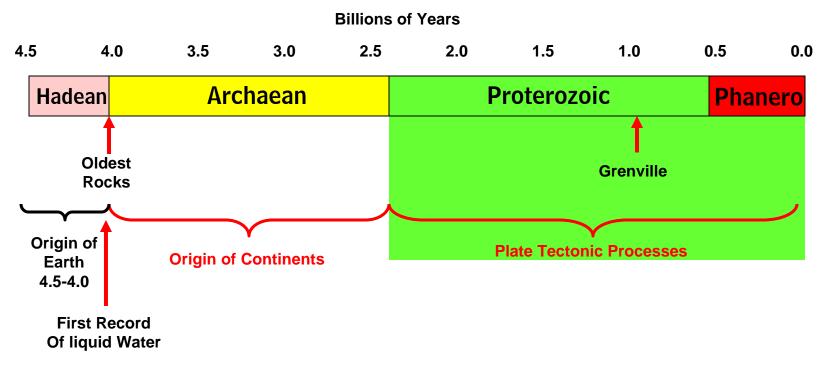
Proterozoic Evolution of North America

JUST HOW LONG IS EARTH HISTORY ? AND WHEN DID THE IMPORTANT THINGS HAPPEN ?

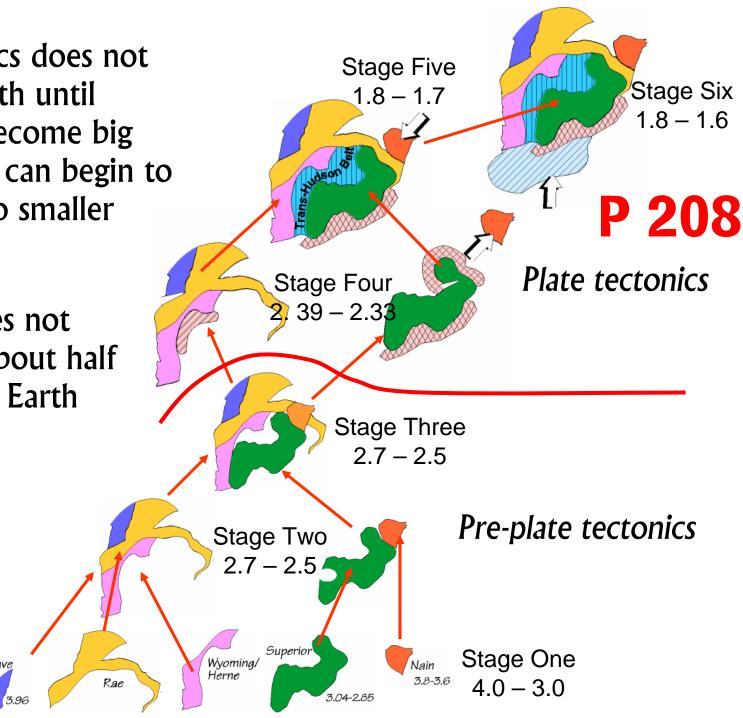


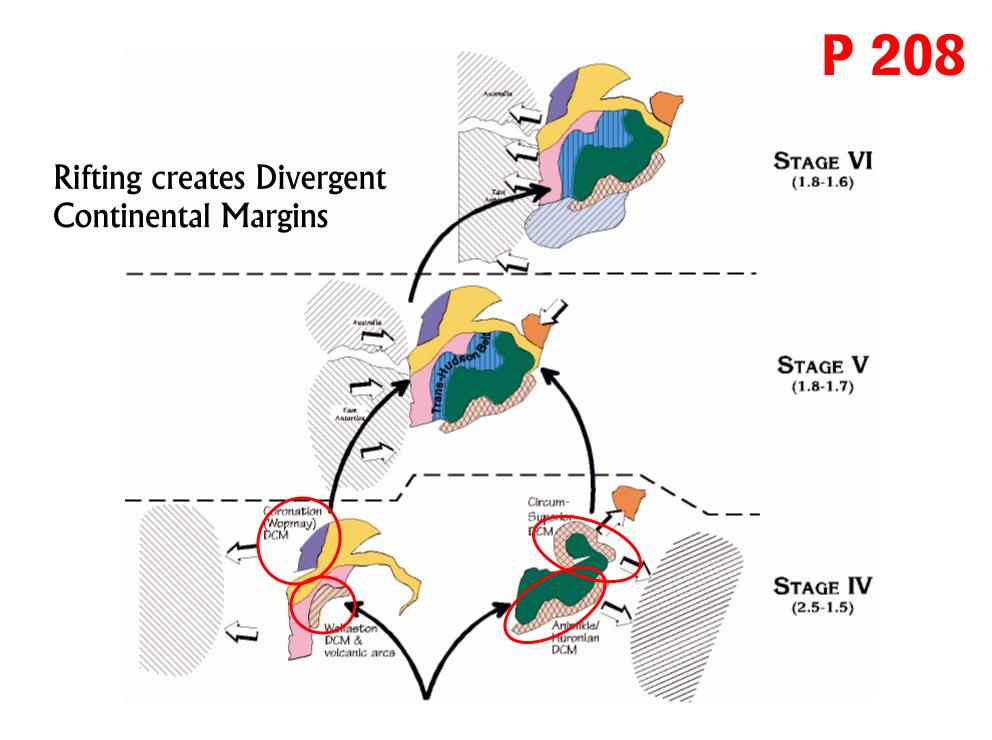
Archaean tectonic evolution of the Earth involved volcanic arcs, evolving into proto-continents, evolving into micro-continents.

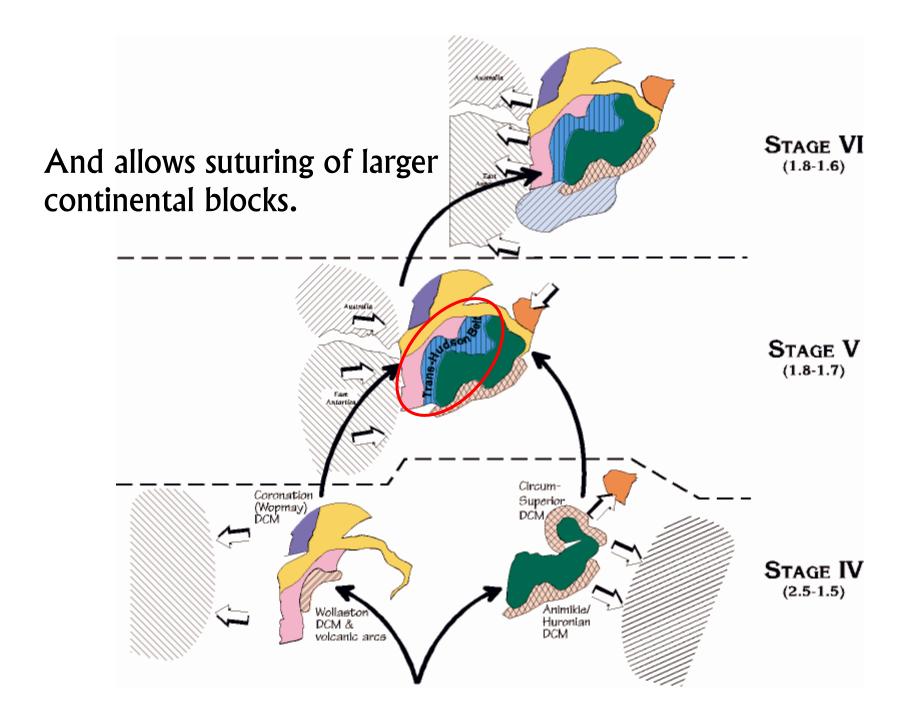
Proterozoic evolution involved the building of the continents, and establishment of supercontinent cycles through plate tectonic mechanisms. Plate tectonics does not begin on Earth until continents become big enough they can begin to rift apart into smaller continents.

And this does not begin until about half way through Earth history.

Slave







The history of the Earth embodies two organizing principles.

1. It is fractal . . .

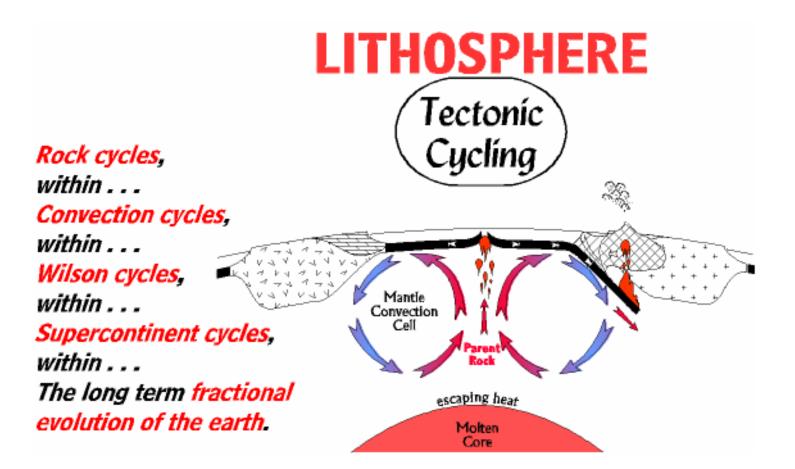
There are patterns, within patterns, within patterns, and cycles, within cycles, within cycles. . . ad infinitum – at all scales of observation.

2. It is evolutionary . . .

As energy continuously passes through the various Earth systems they undergo continuous change, ...

...generally becoming more complex with time through elaboration, fractionation, and self organization.

Cycles Within Cycles The fractal nature of the Earth

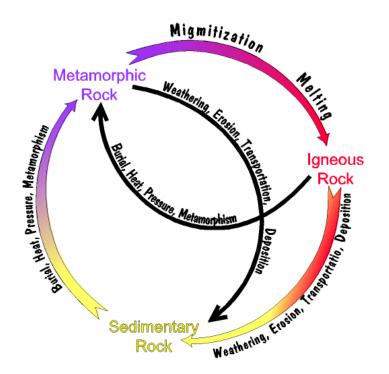


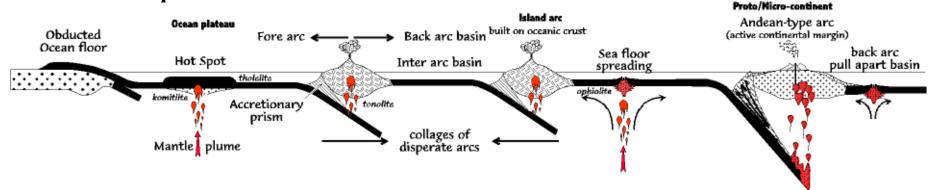
Cycles Within Cycles

ROCK CYCLES . . .

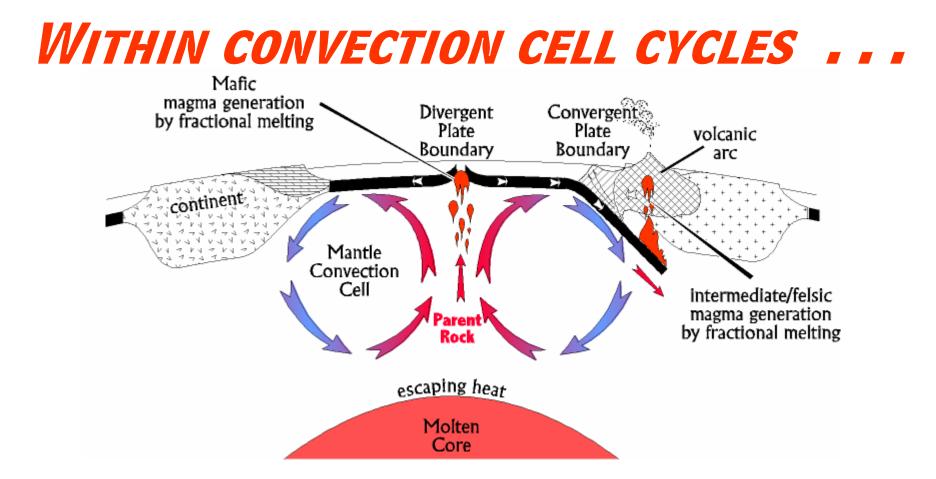
All rocks can be transformed Into other rocks

And with every cycle of subduction the igneous rocks that are generated, weather into sedimentary rocks, some of which become metamorphosed.





Cycles Within Cycles



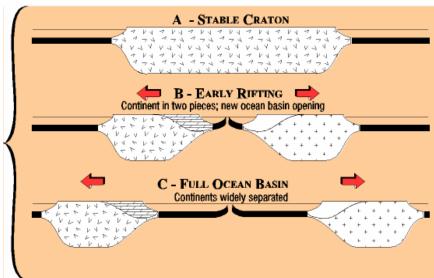
The rock cycle is driven by tectonic energy from the molten cores escaping inexorably to space.

Cycles Within Cycles P 203

WITHIN WILSON CYCLES . . . Wilson Closing Phase

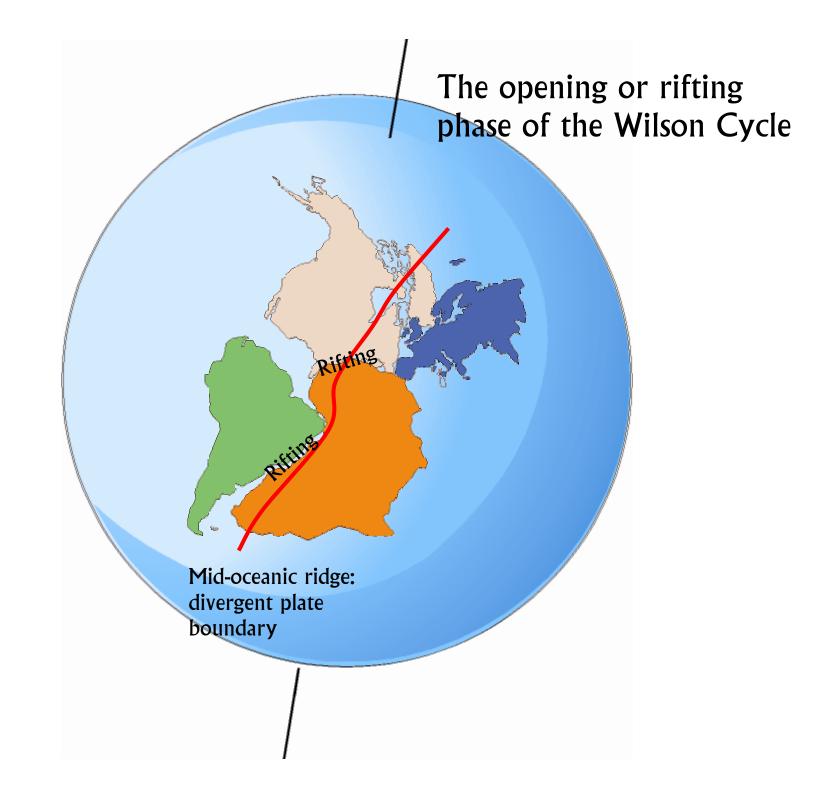
The opening and closing of an ocean basin

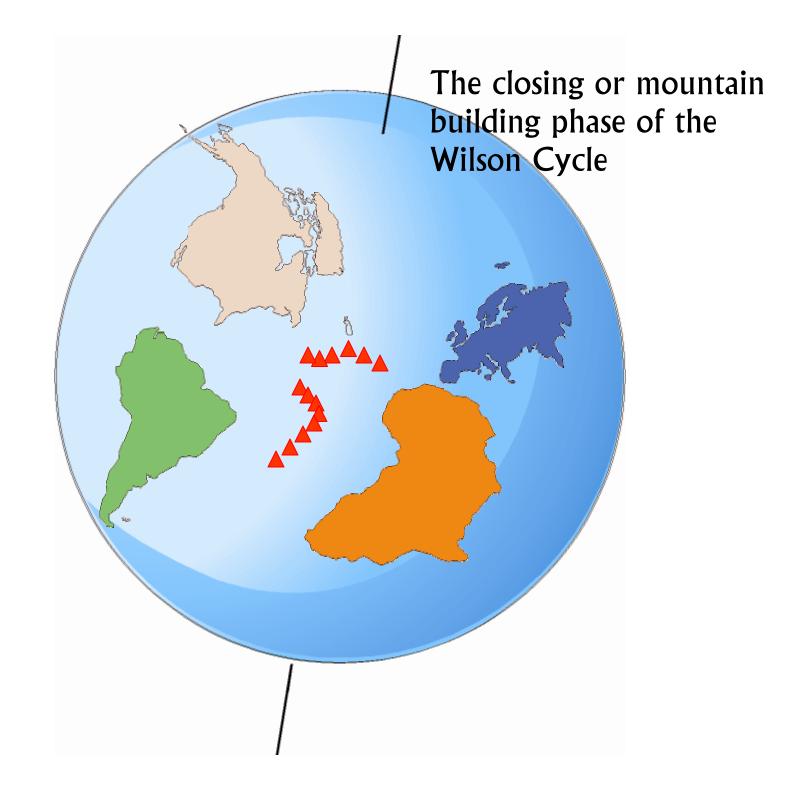
Wilson Opening Phase



Leading to . . .

D - SUBDUCTION ZONE Ocean basin begins to close E - CLOSING REMNANT Ocean basin almost closed: collision about to occur OCEAN BASIN F - COLLISION OROGENGY Right continent overrides left continent; ocean basin closed **G** - PENEPLAINED MOUNTAIN Mountains eroded to sea level tectonic stability agair

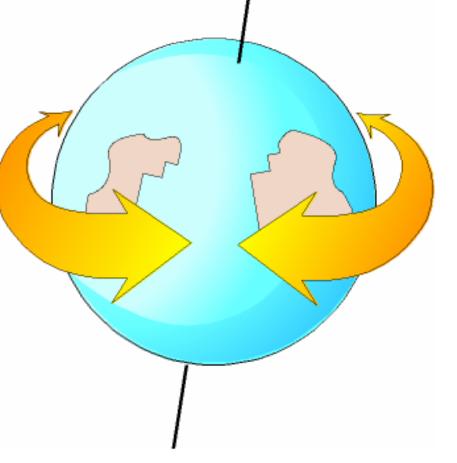




Cycles Within Cycles

WITHIN SUPERCONTINENT CYCLES . . .

In a supercontinent model continents sweep back and forth across the globe. Simplistically the result would be two Wilson cycles, each half out of phase with each other. When a supercontinent breaks apart on one side of the globe, initiating an opening phase, it also begins a closing half of the other cycle on the other side of the globe.

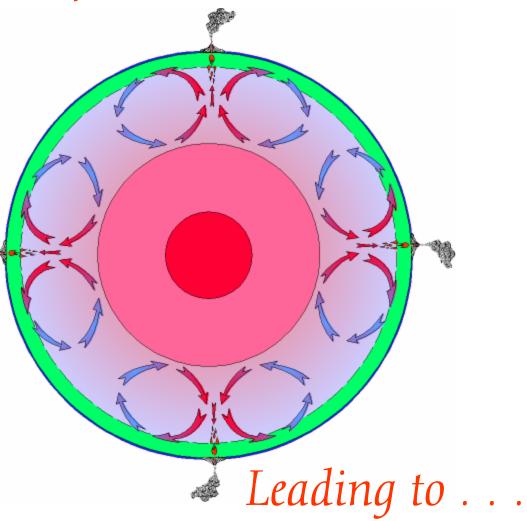


But, more realistic

Cycles Within Cycles

A Complex of Convection Cells

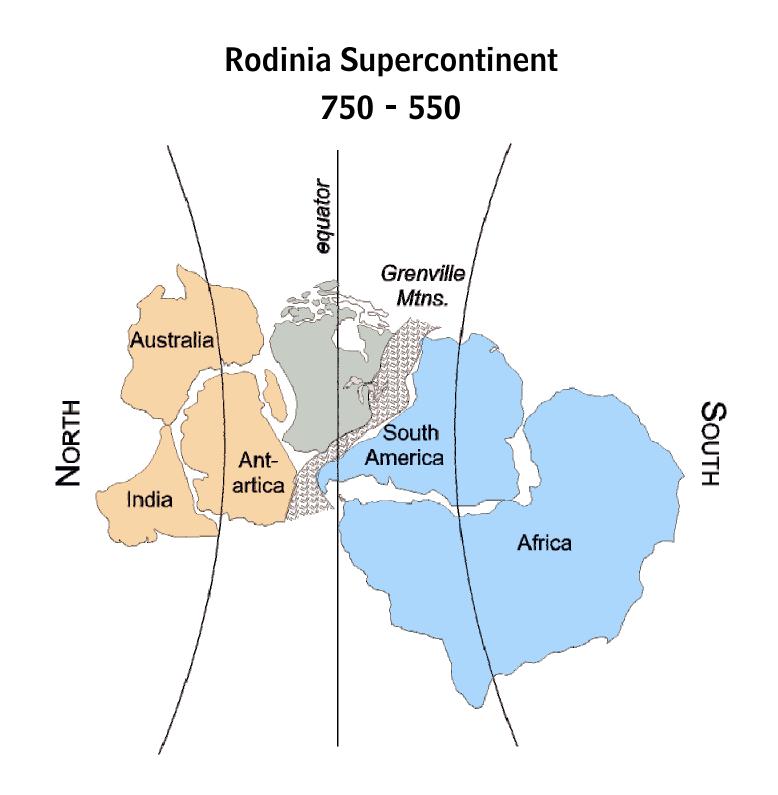
But, the Earth is three dimensional and has a complex system of convection cells making the possible outcomes more complex than simple.

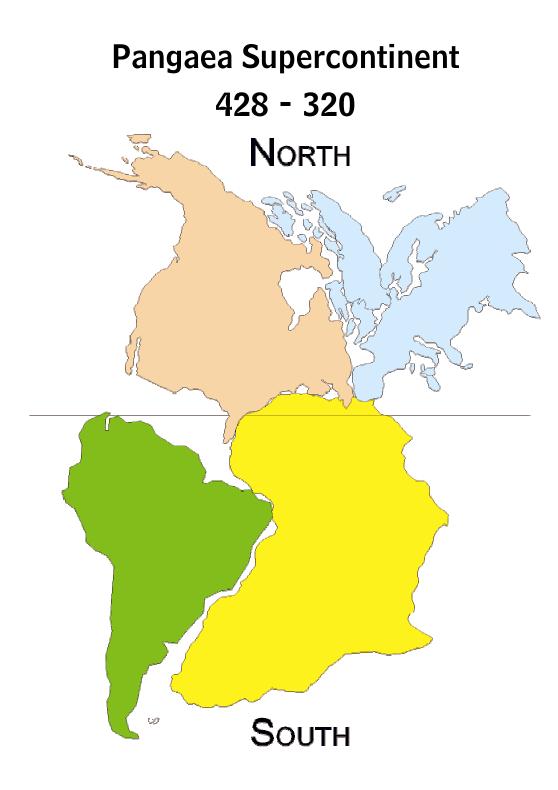


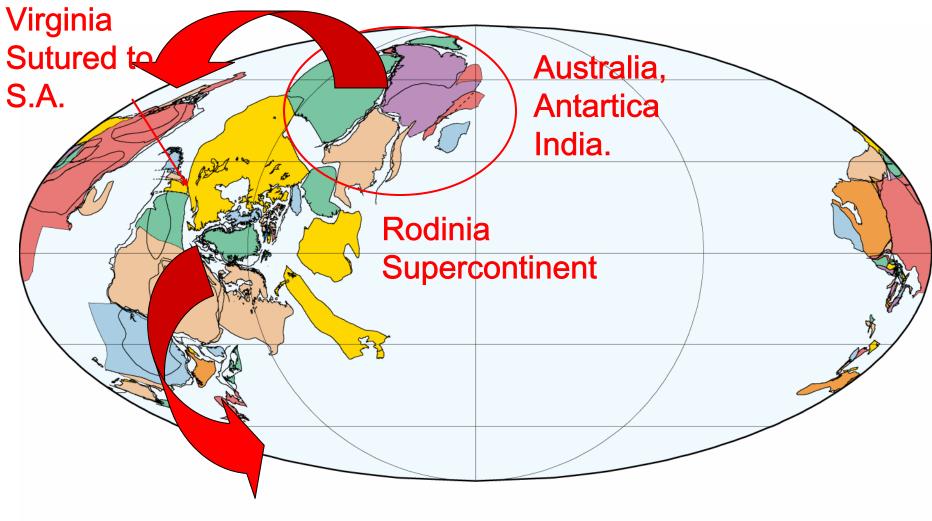


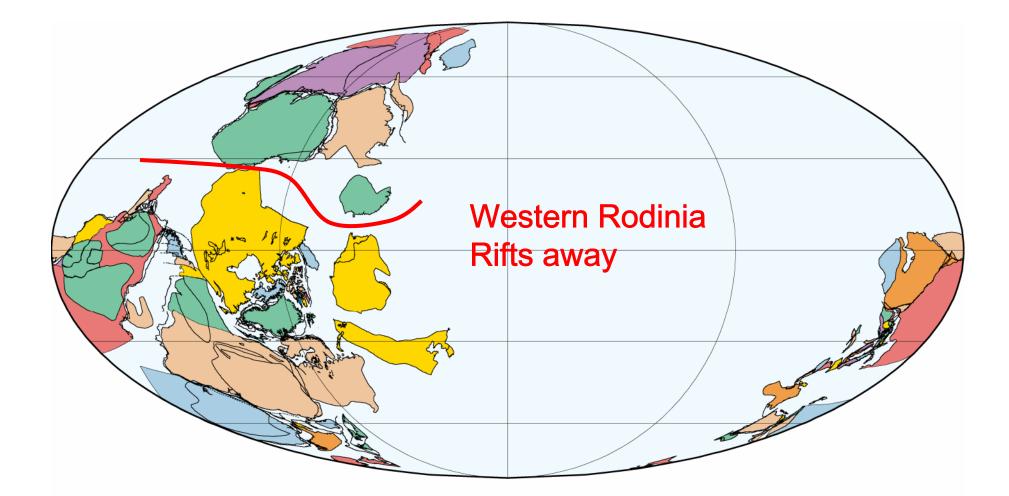
Cycles Within Cycles Evolutionary Rock Cycles And as a result of every cycle the continents get bigger, and more igneous, sedimentary, and metamorphic rocks are generated by fractionation. Multiple Wilson & Supercontinent Cycles Migmitization Metamorphic Build, Heat, Pressure, Metamorphi. Isolated Large Volcanic Continents Arcs Erosion Sedimentary Weathering. Rock Rock and Earth Evolution

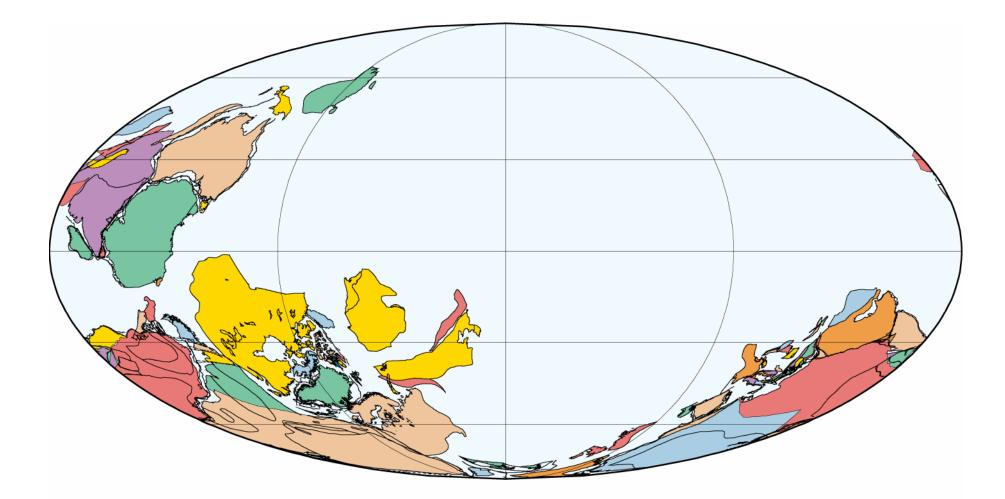
Supercontinent Cycles

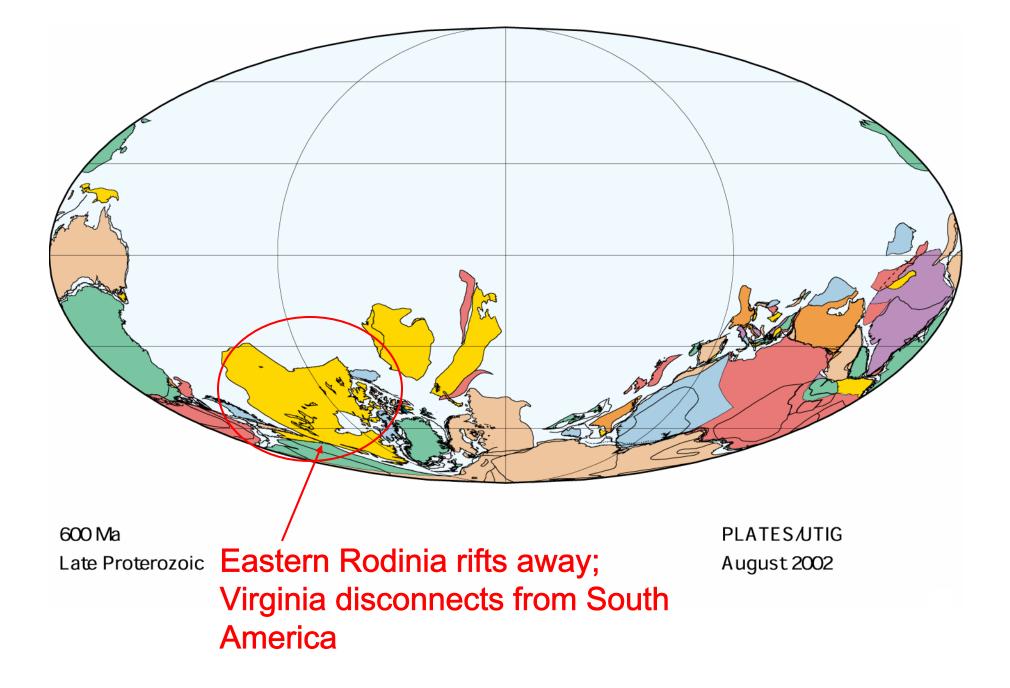


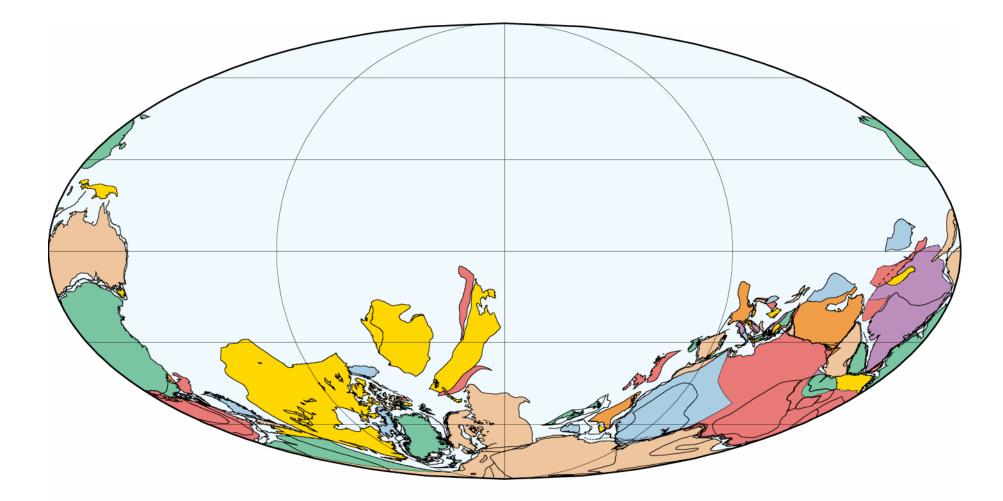


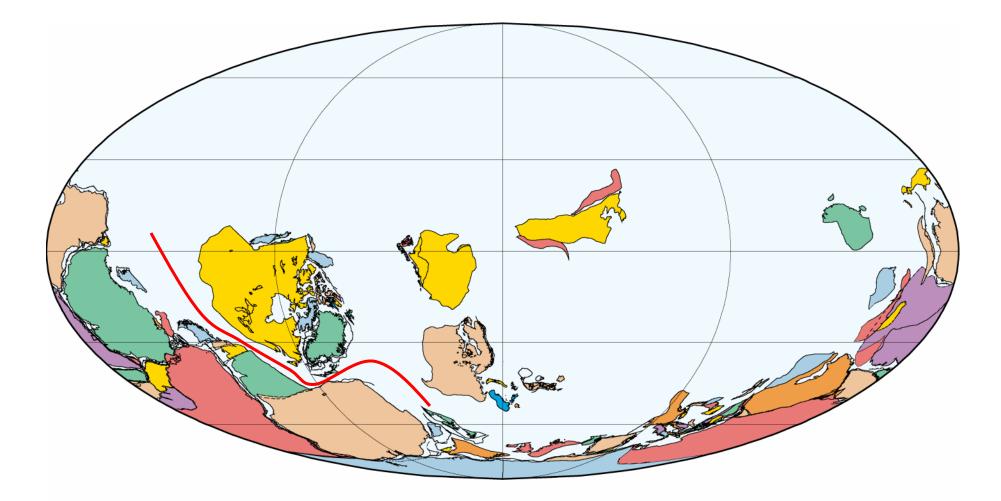




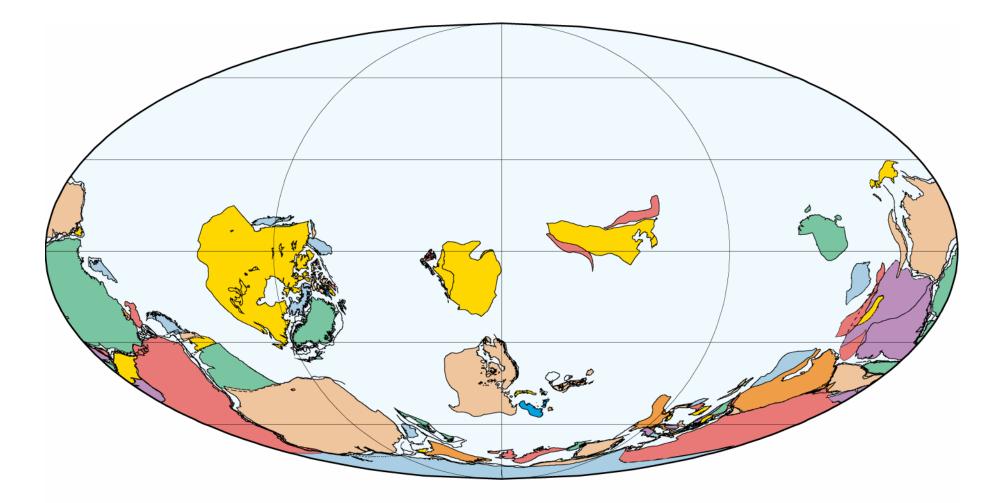




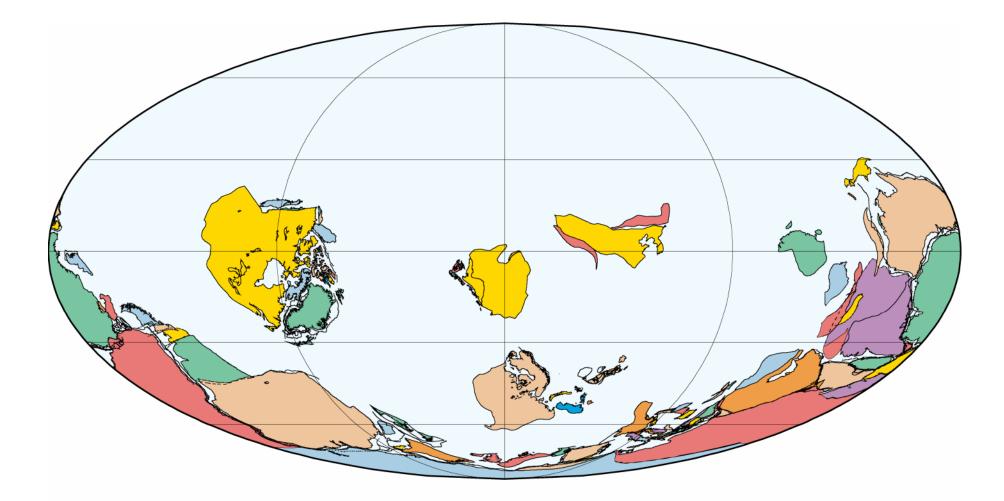




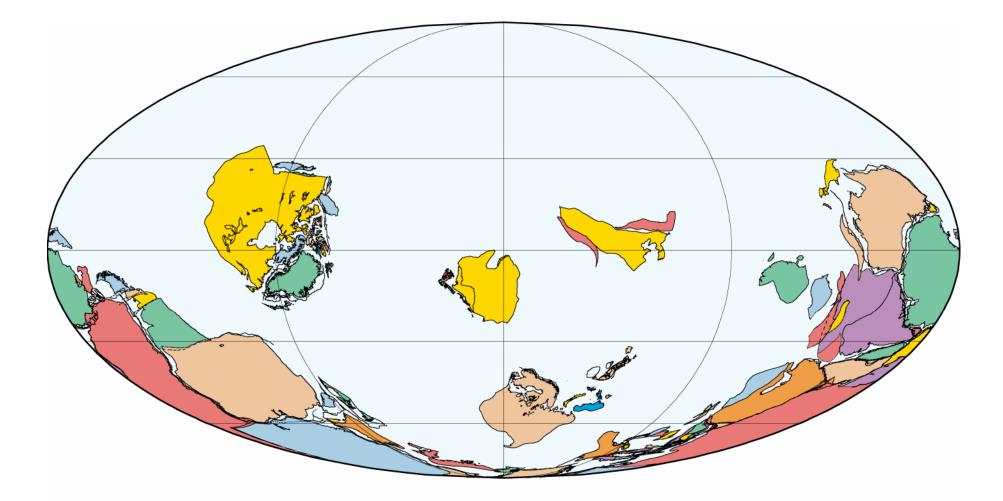
540 Ma Nemakitian-Daldynian (Early Cambrian)



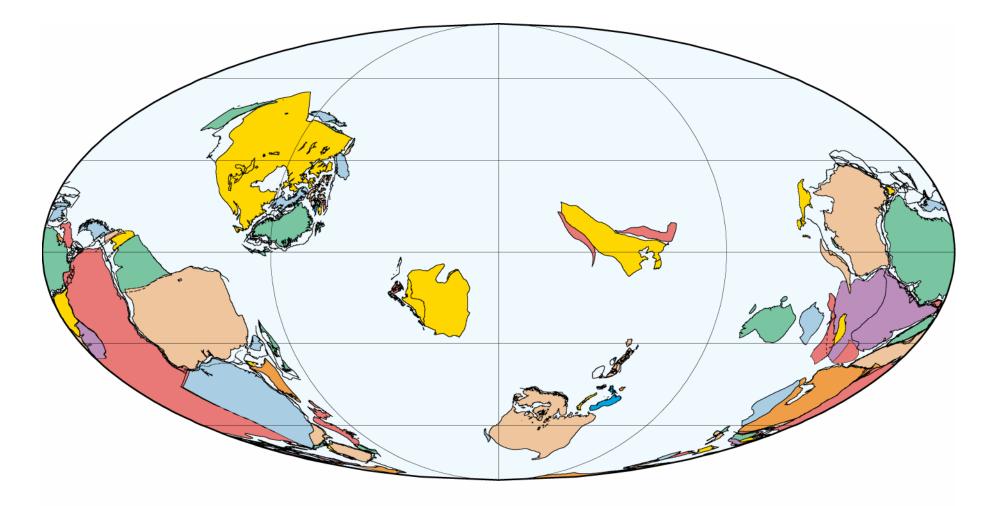
530 Ma Late Tommotian/Early Atdabanian (Early Cambrian)



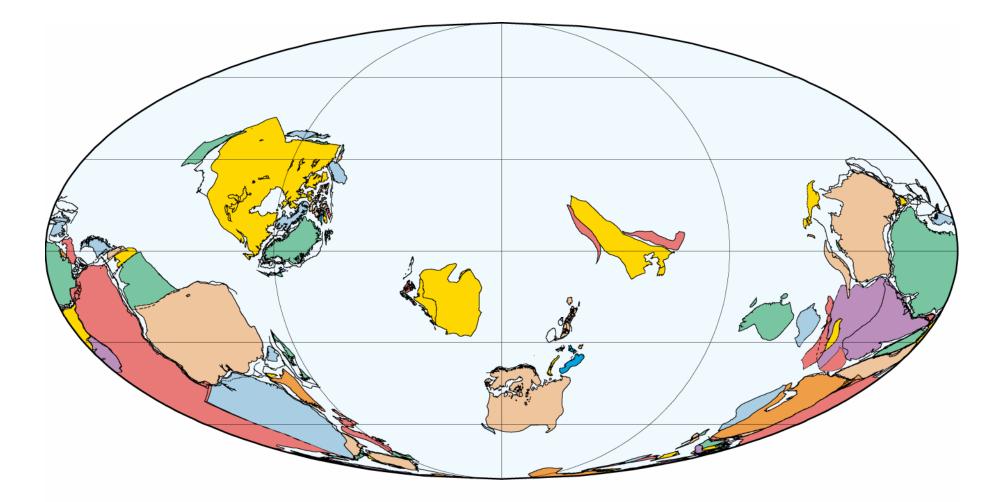
520 Ma Lenian (Early Cambrian)



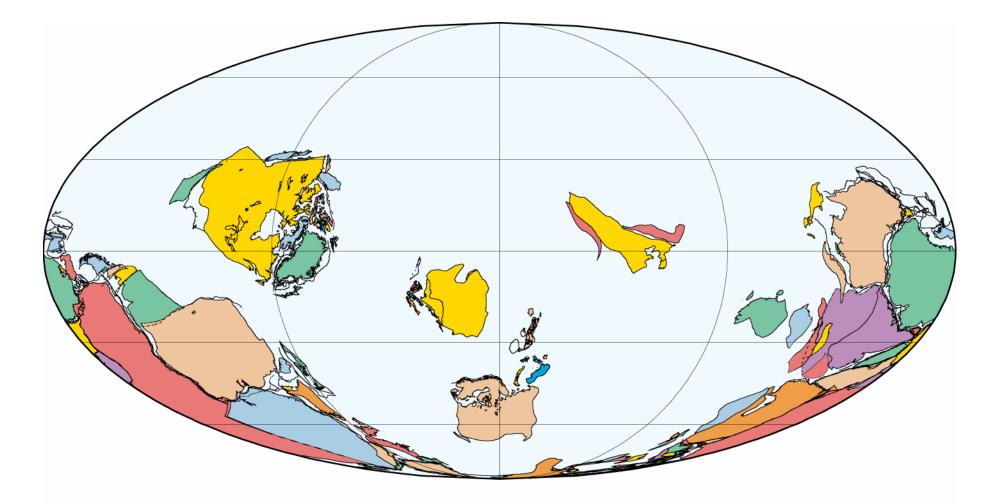
510 Ma Middle Cambrian



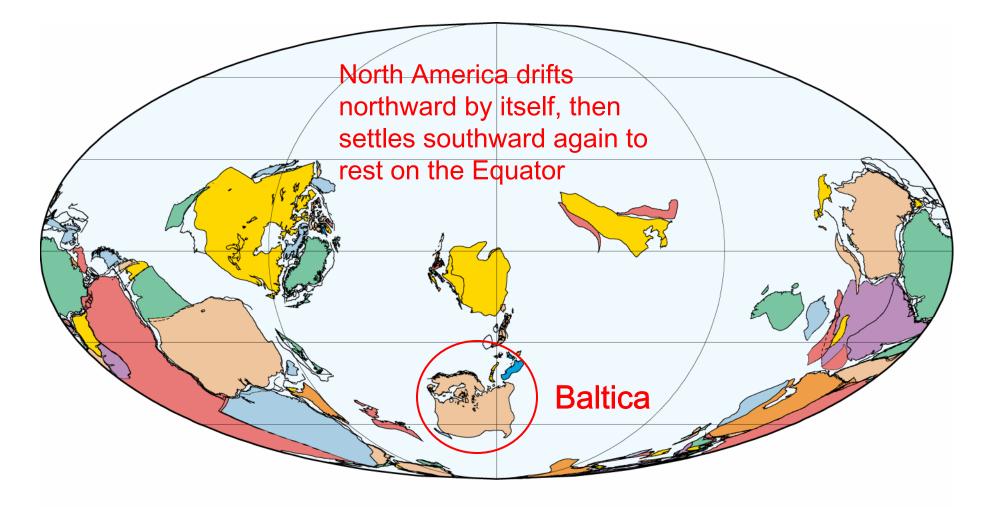
500 Ma Late Cambrian



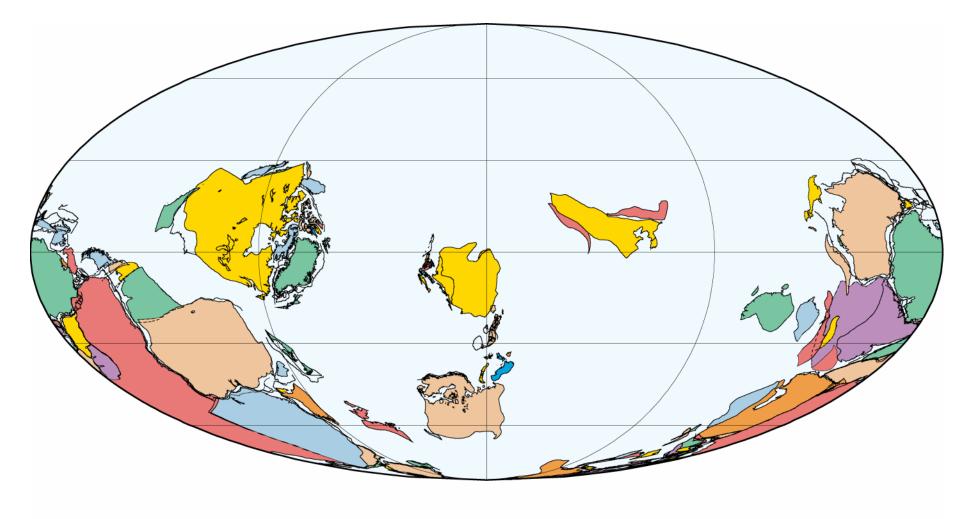
490 Ma Tremadocian (Early Ordovician)



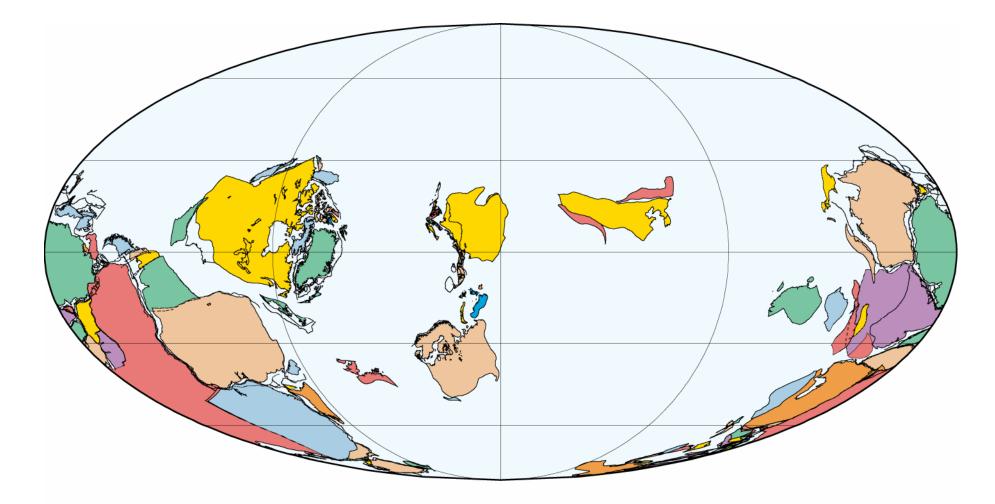
480 Ma Arenigian (Early Ordovician)



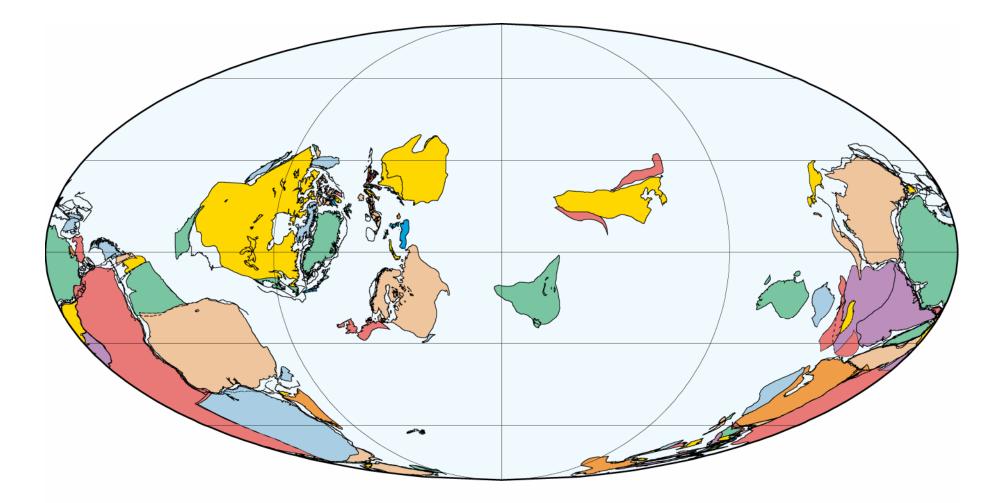
470 Ma Late Arenigian/Early Llanvirnian (Early/Middle Ordovician)



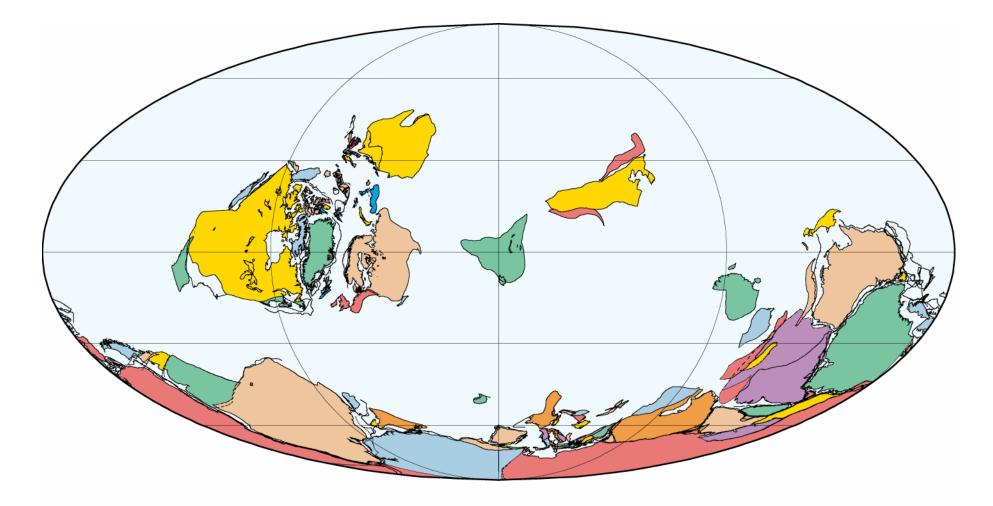
470 Ma Late Arenigian/Early Llanvirnian (Early/Middle Ordovician)



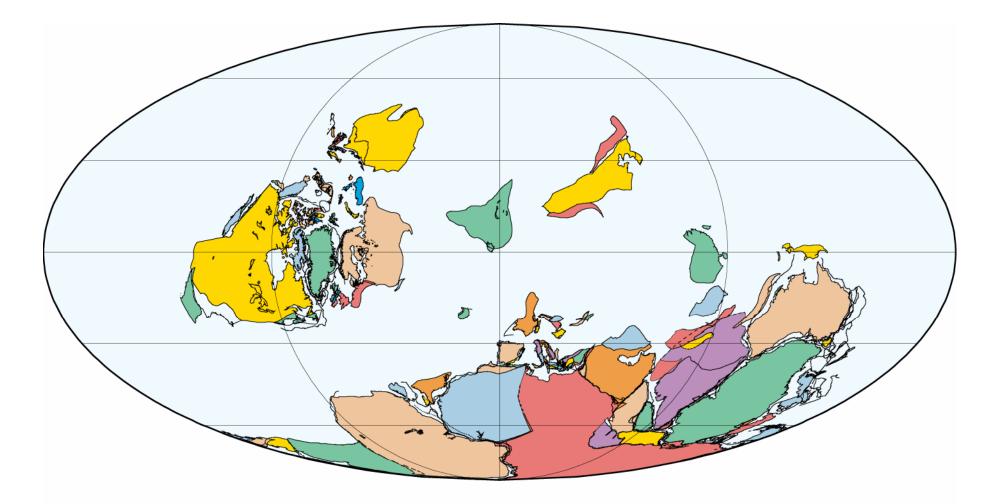
460 Ma Llandeilan (Middle Ordovician)



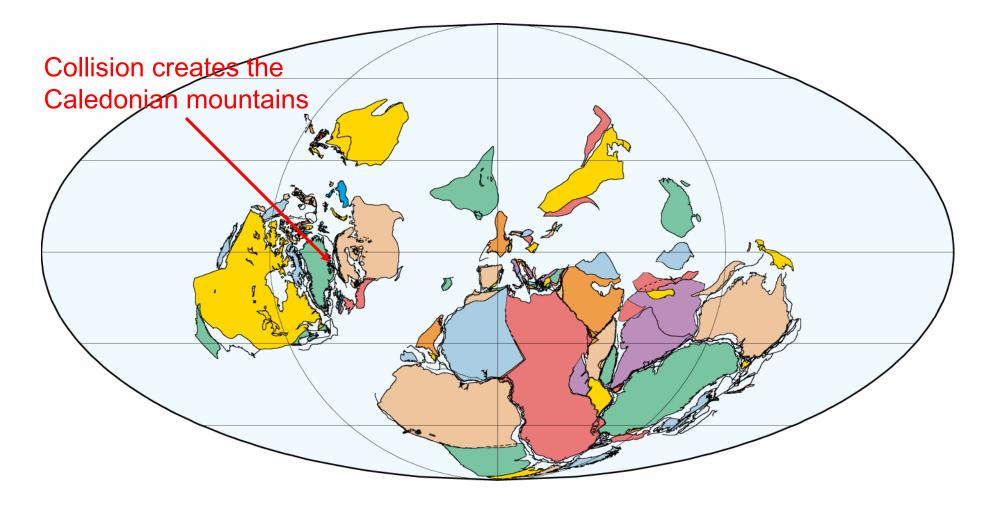
450 Ma Caradocian (Late Ordovician)



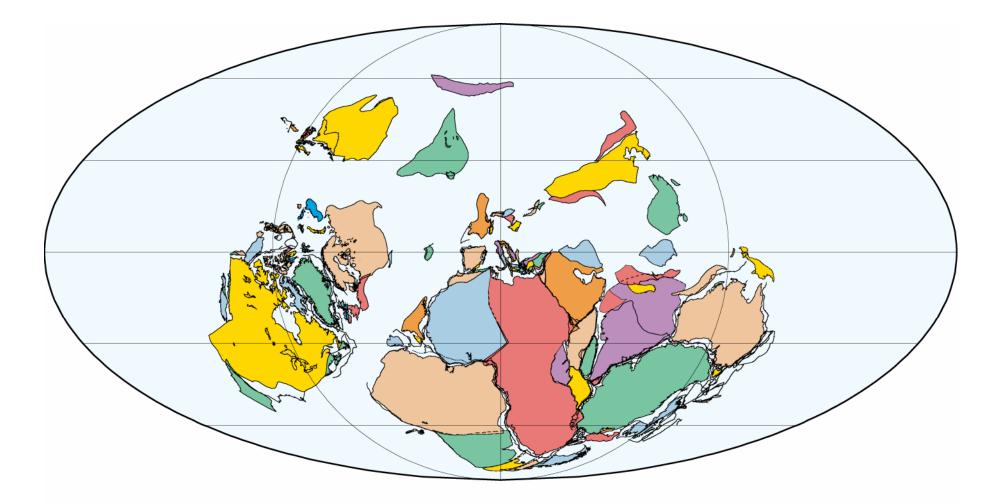
440 Ma Early Llandoverian (Early Silurian)



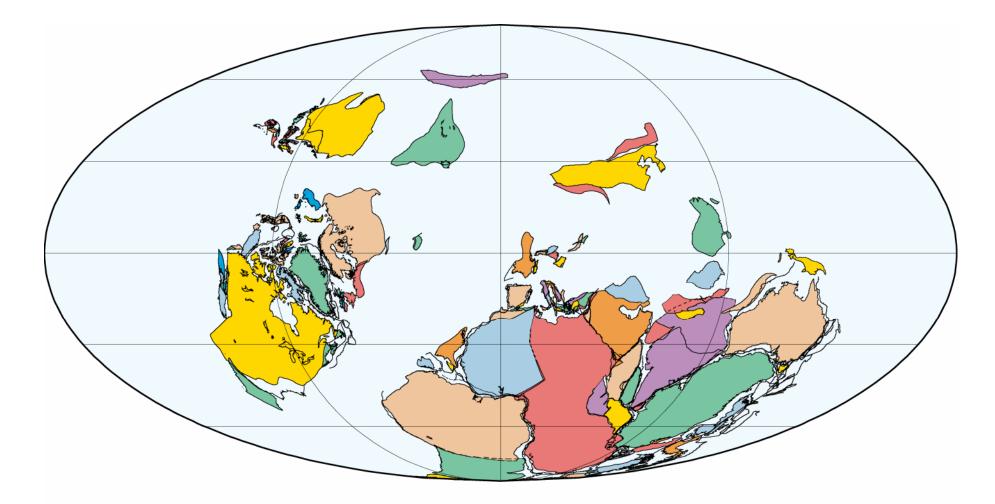
430 Ma Late Llandoverian (Early Silurian)



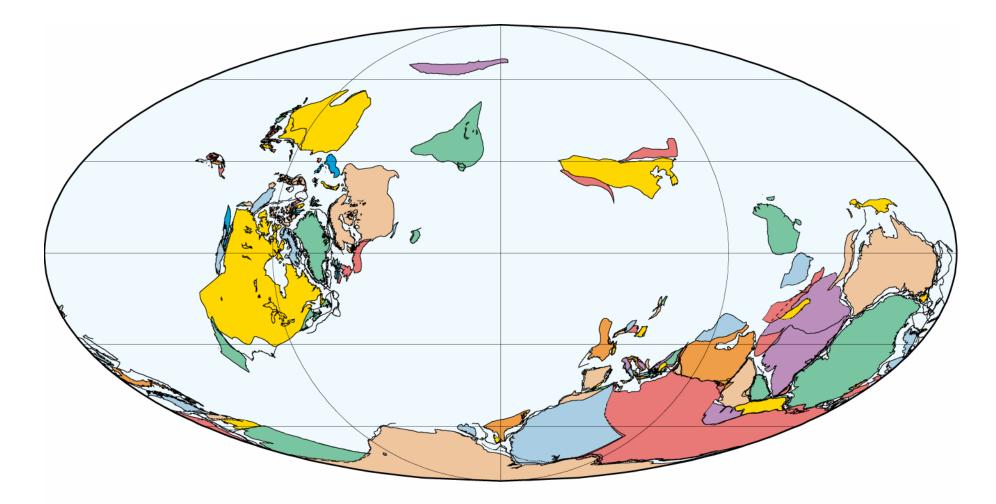
420 Ma Ludlovian (Late Silurian)



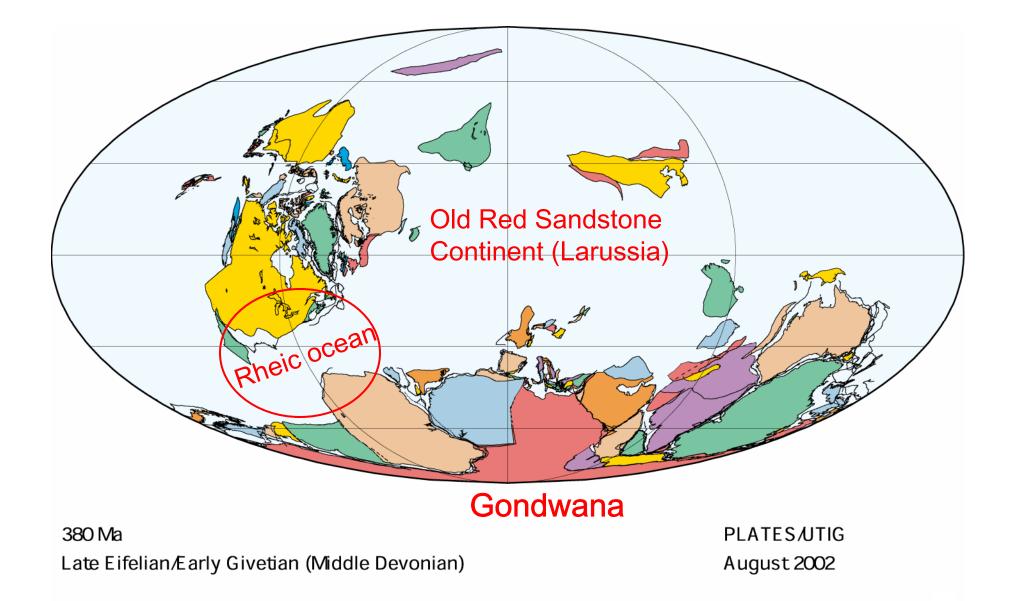
410 Ma Early Praghian (Early Devonian)

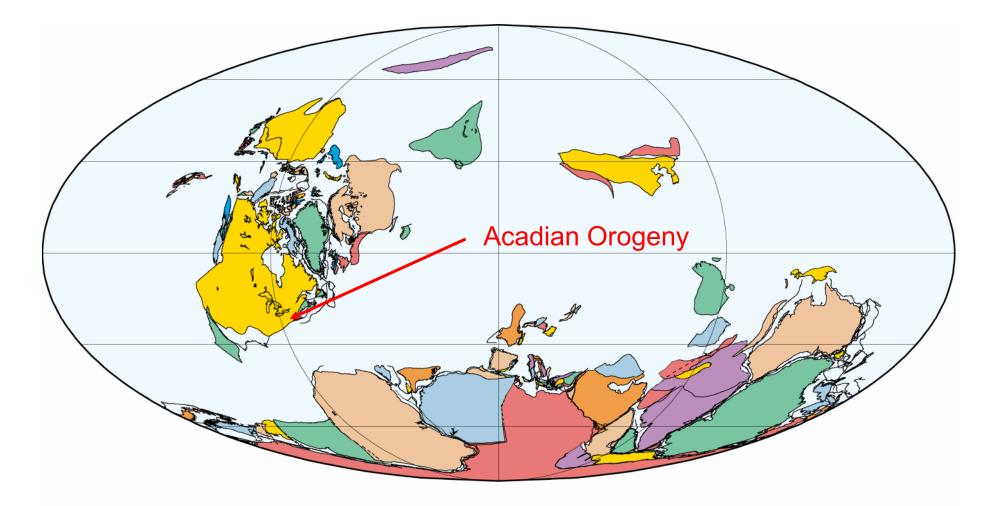


400 Ma Late Praghian/Early Emsian (Early Devonian)



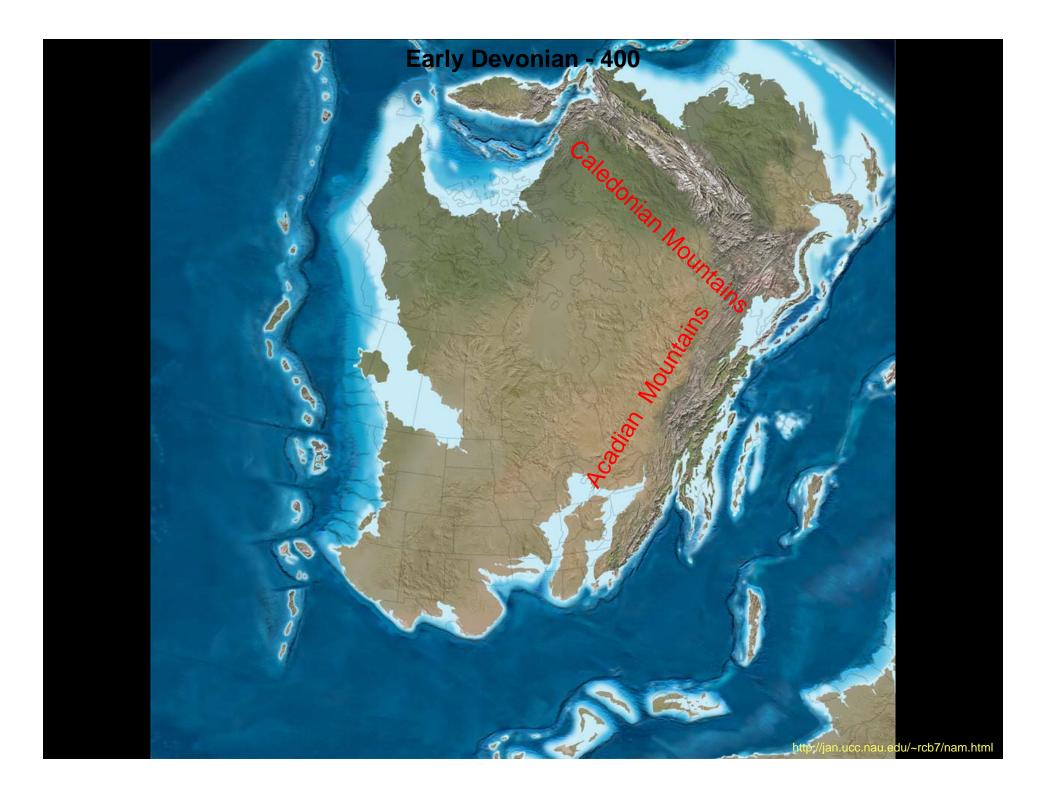
390 Ma Early Eifelian (Early Devonian)

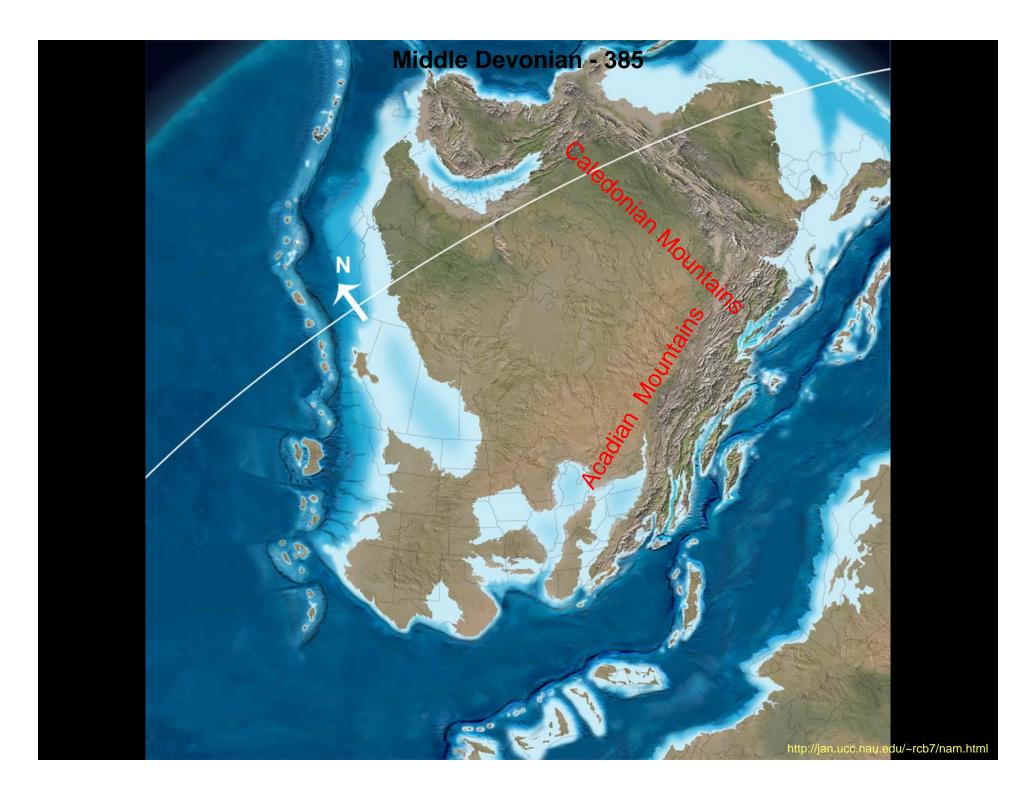


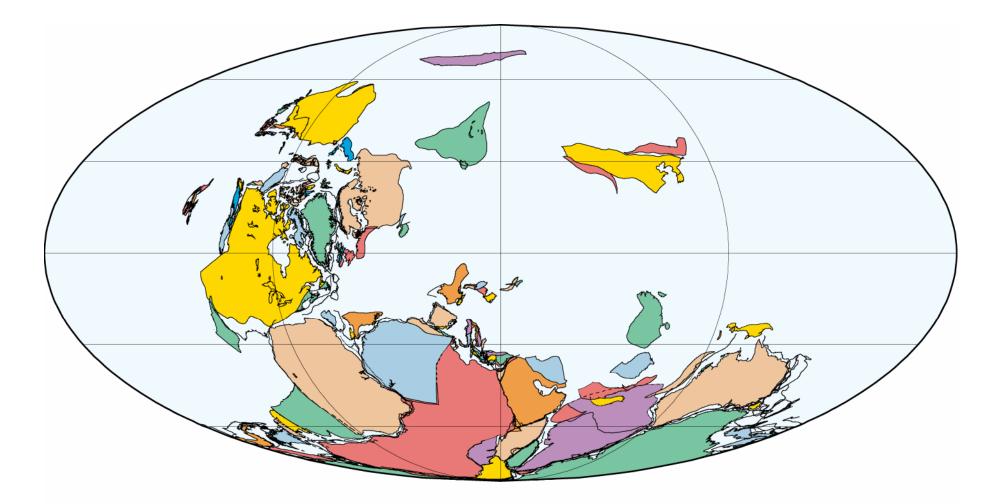


380 Ma Late Eifelian /Early Givetian (Middle Devonian)

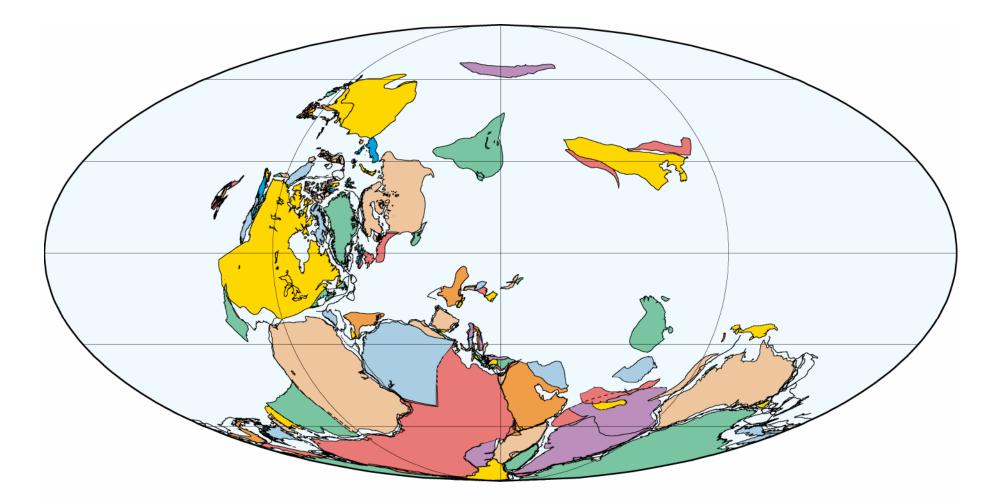




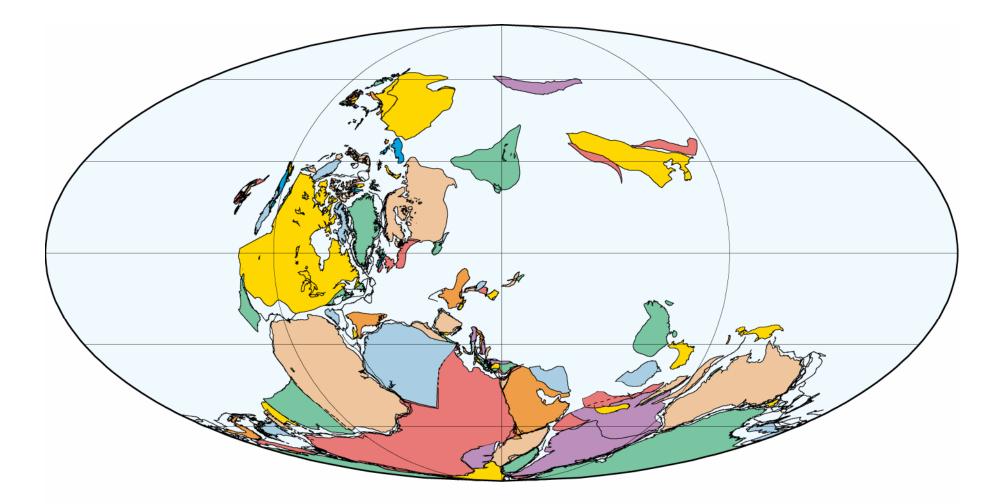




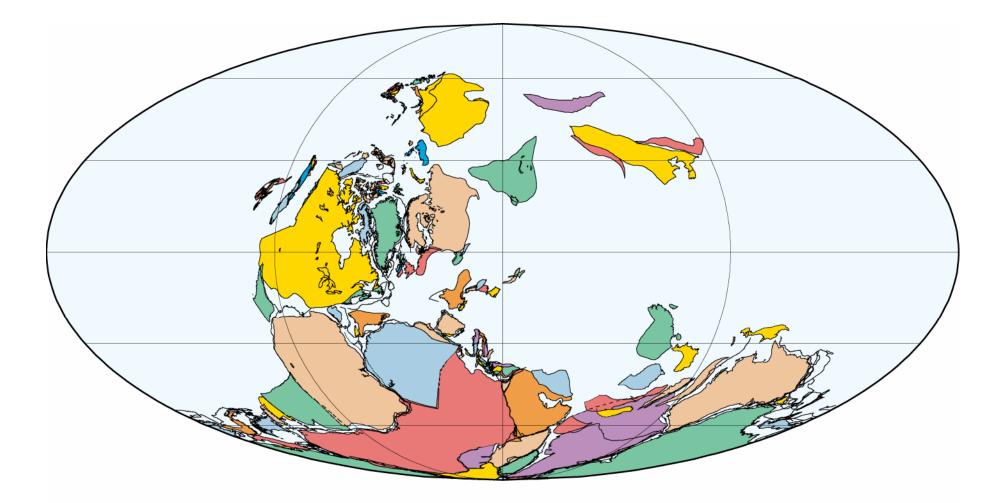
370 Ma Late Givetian/Early Frasnian (Late Devonian)



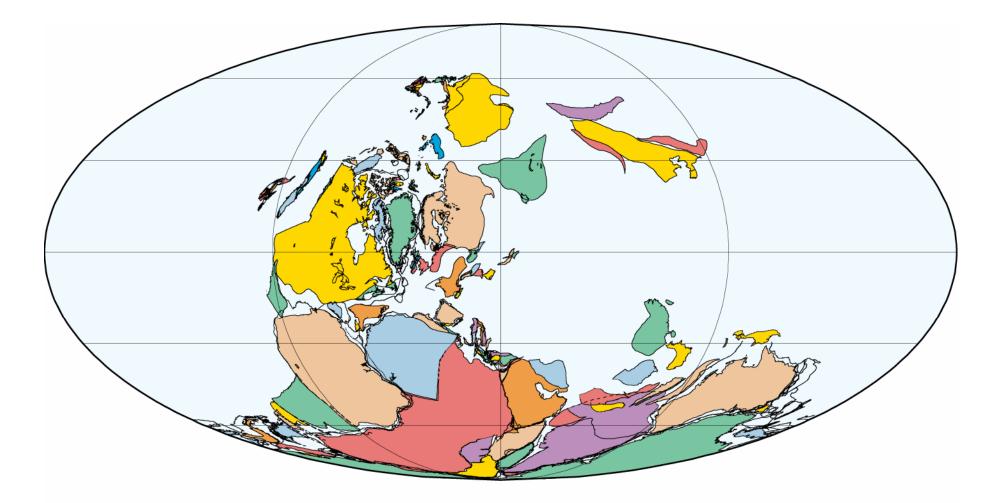
360 Ma Famennian (Late Devonian)



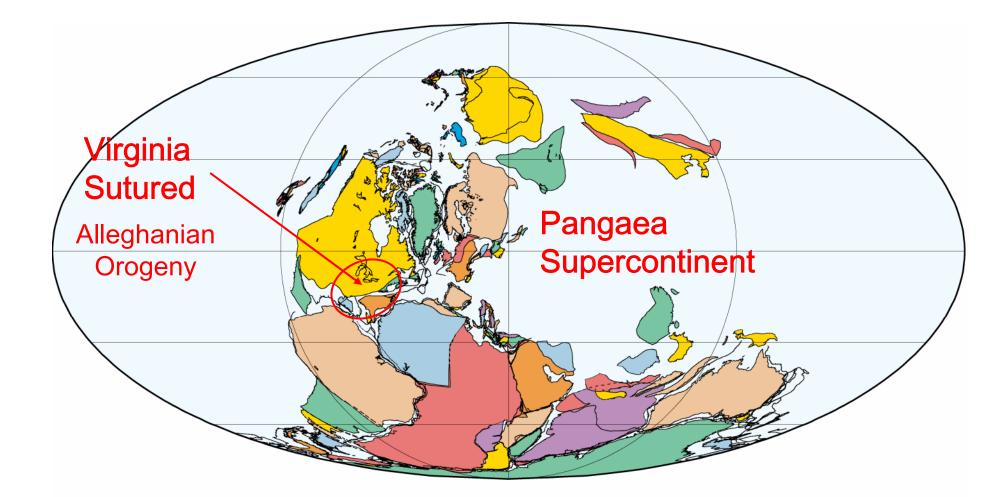
350 Ma Tournaisian (Mississippian)



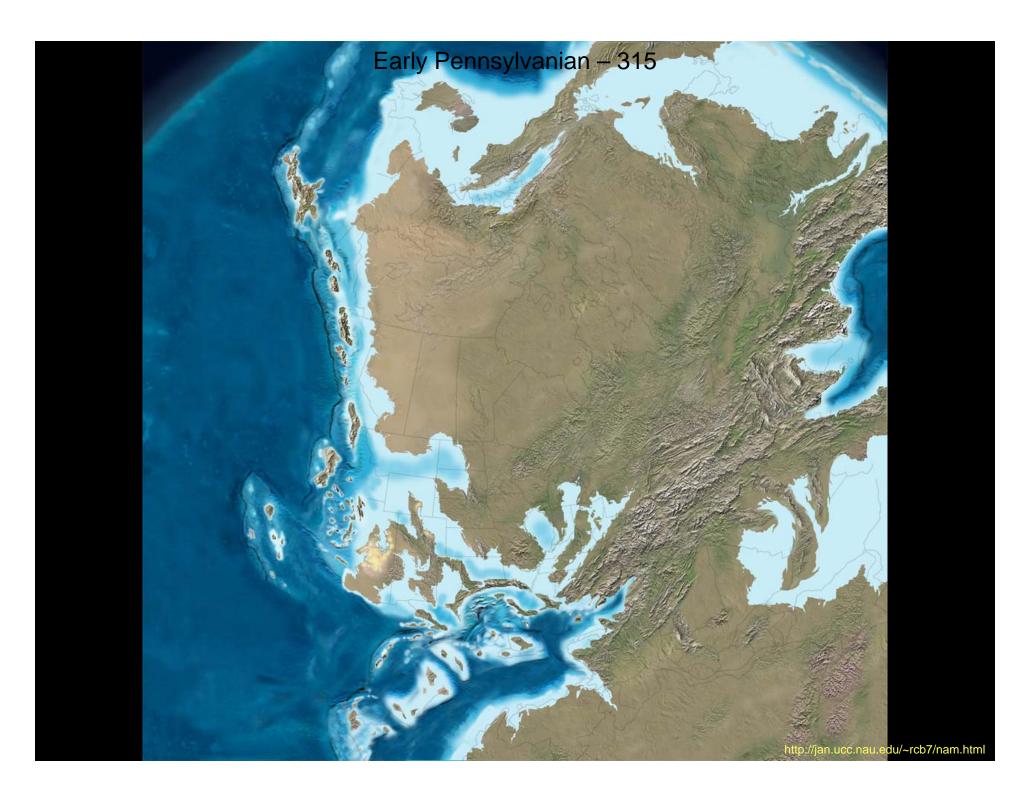
340 Ma Early Visean (Mississippian)

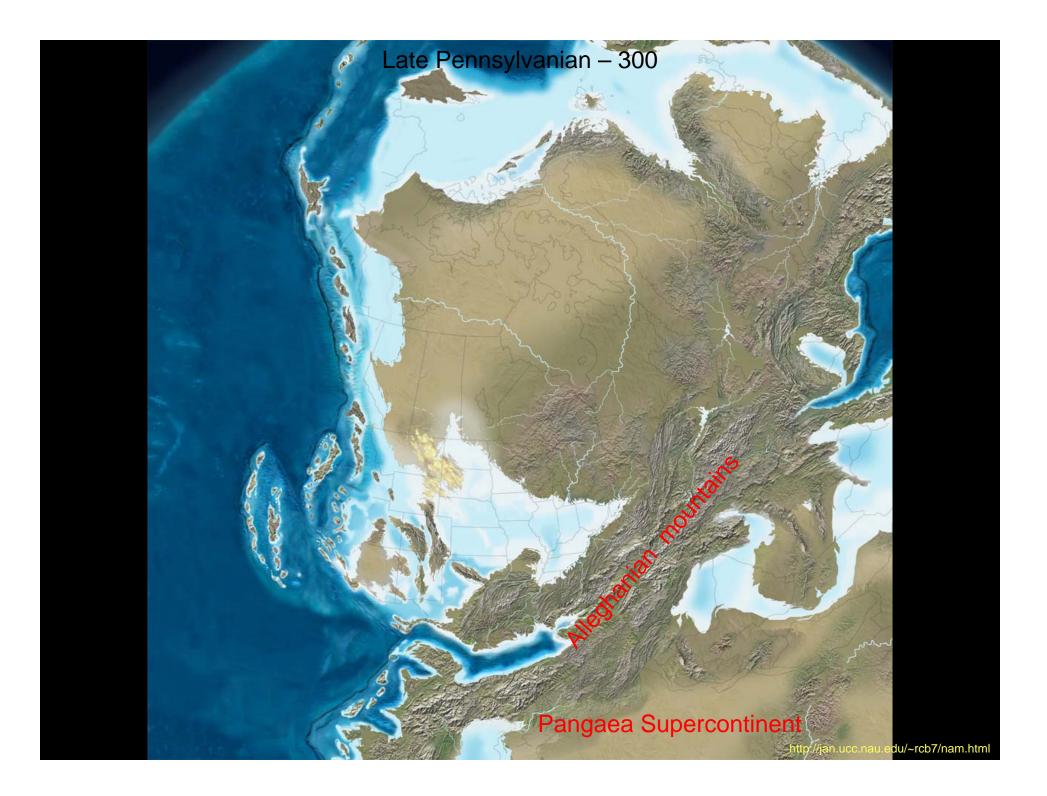


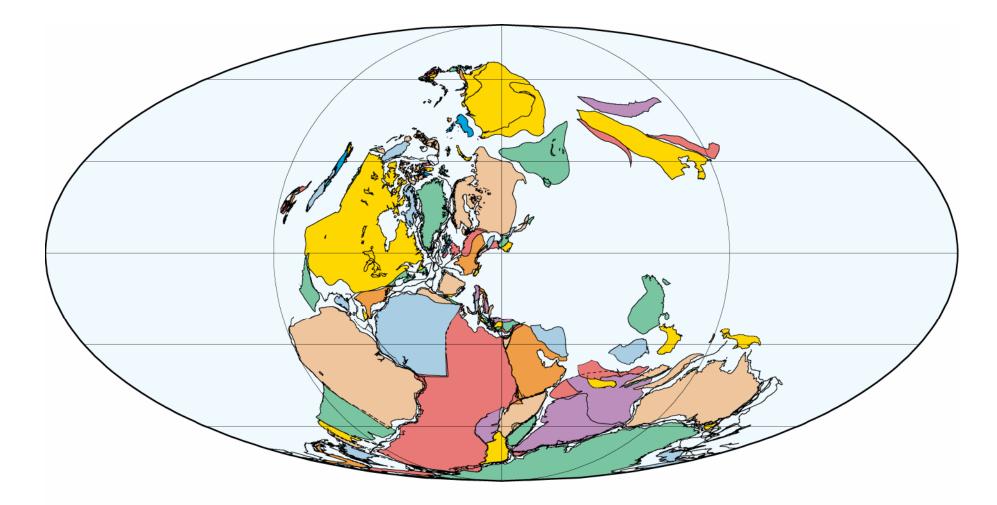
330 Ma Late Visean (Mississippian)



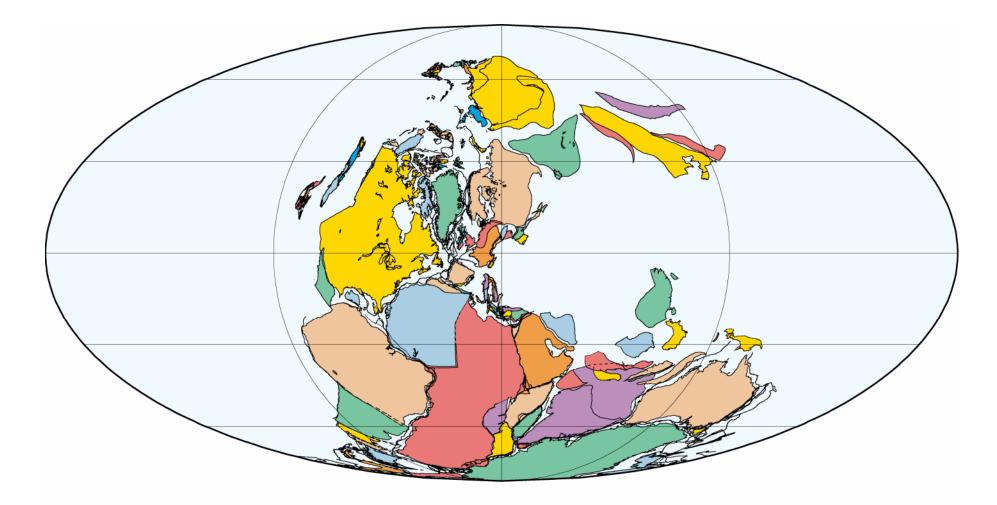
320 Ma Bashkirian (Pennsylvanian)



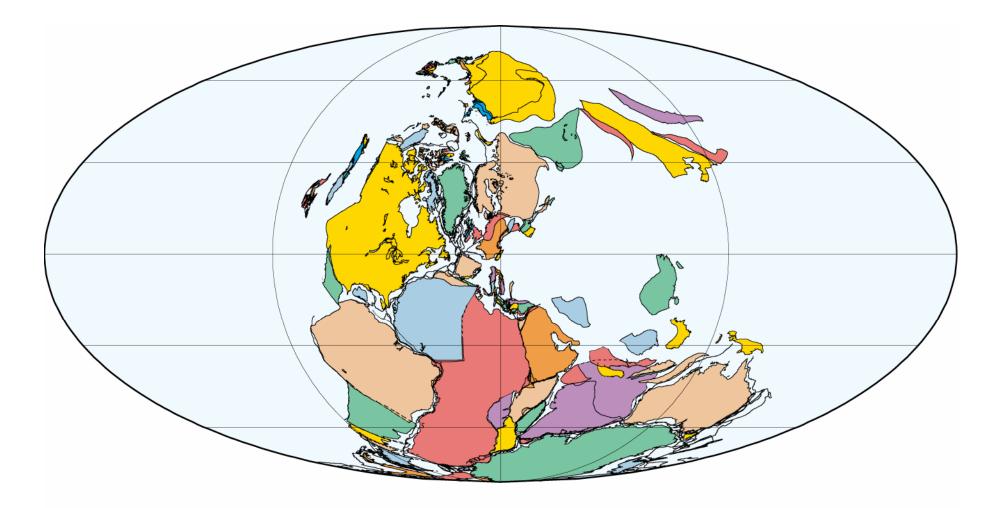




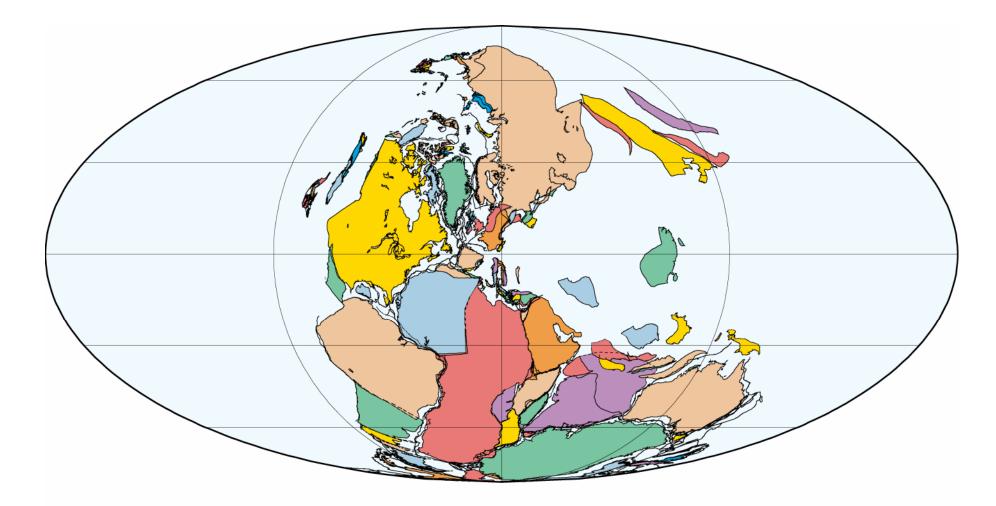
310 Ma Moscovian (Pennsylvanian)



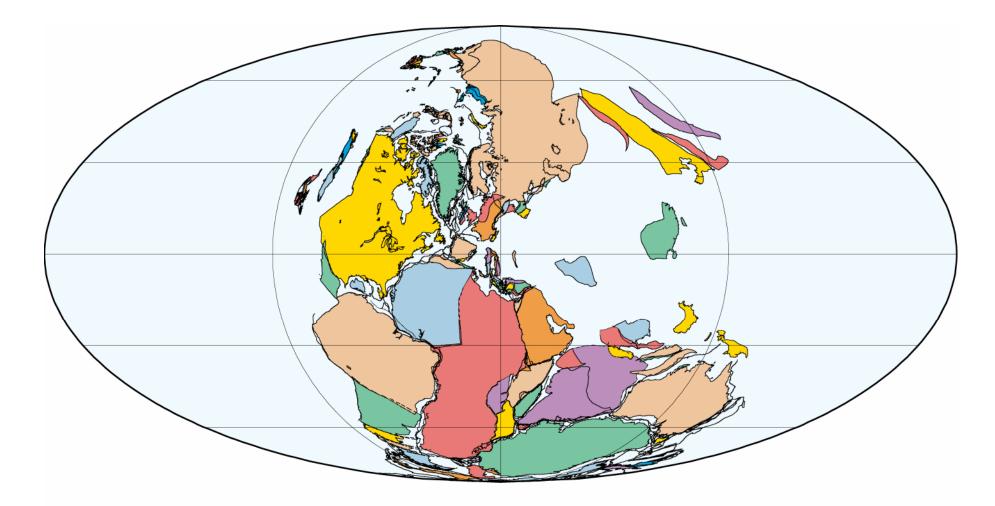
300 Ma Kasimovian (Pennsylvanian)



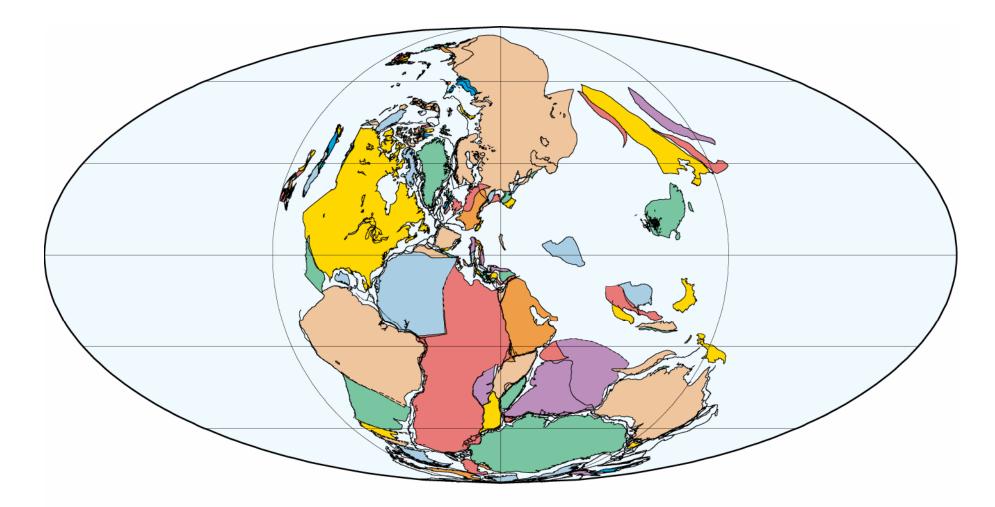
290 Ma Late Gzelian/Early Asselian (Pennsylvanian/Permian)



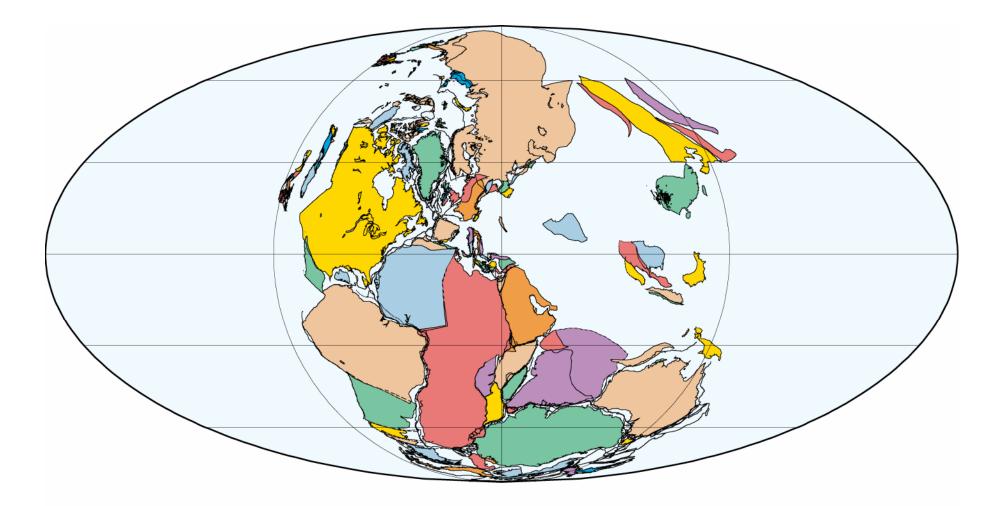
280 Ma Early Sakmarian (Early Permian)



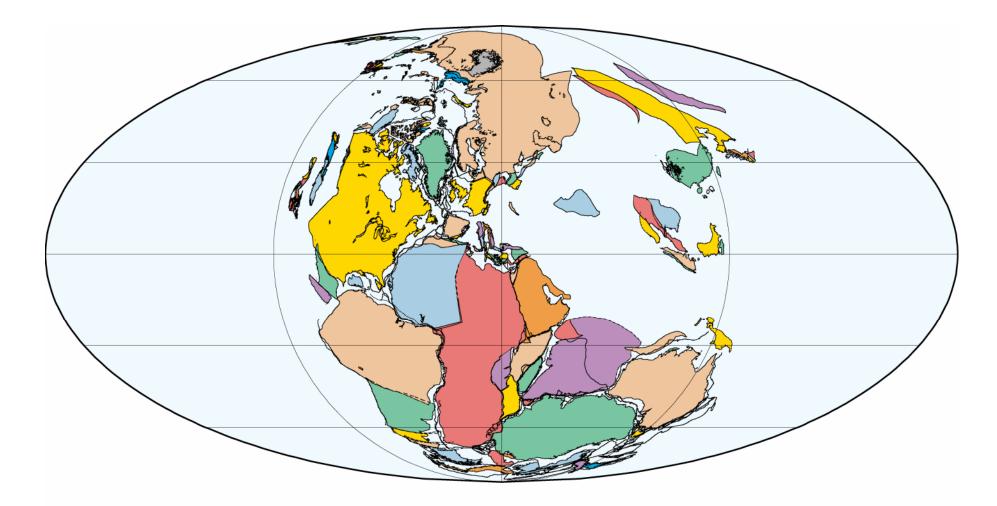
270 Ma Late Sakmarian (Early Permian)



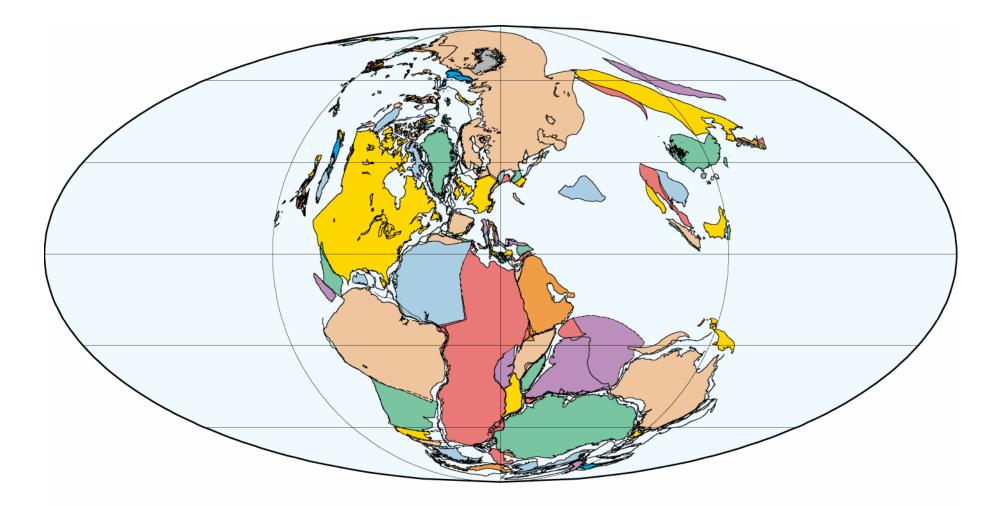
260 Ma Late Artinskian/Early Kungurian (Early Permian)



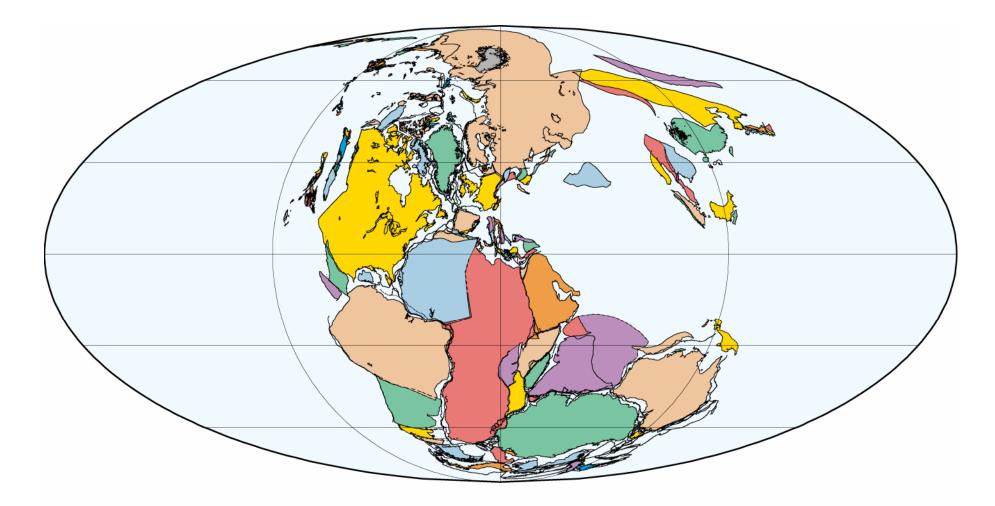
250 Ma Tatarian (Late Permian)



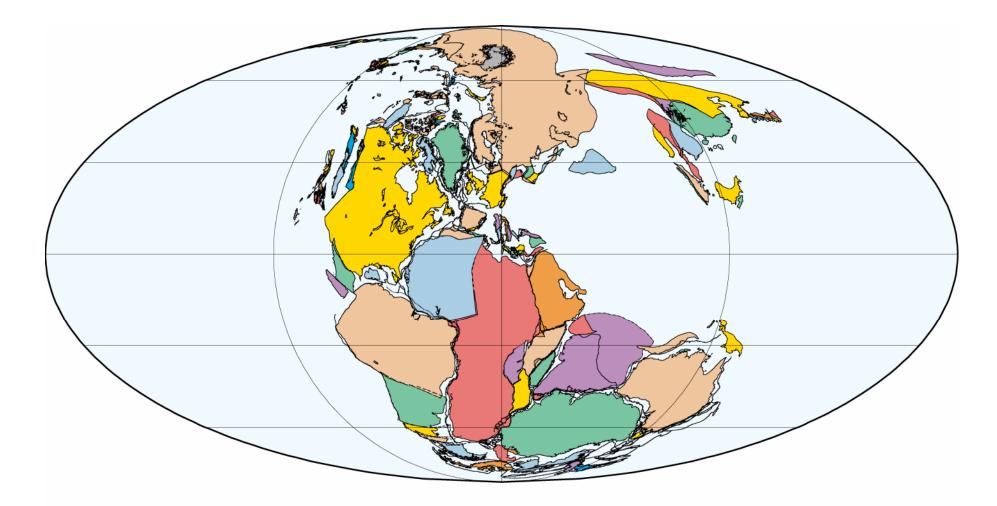
240 Ma Anisian (Middle Triassic)



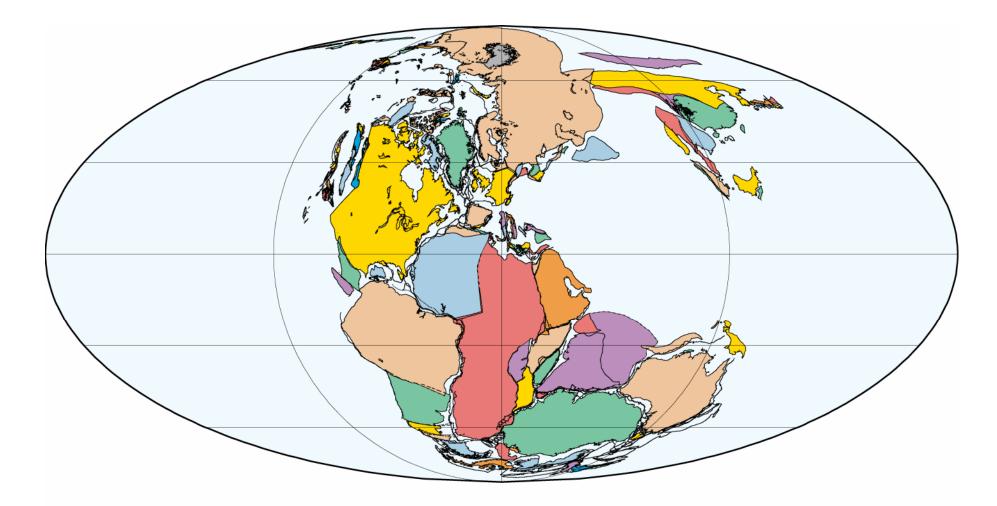
230 Ma Ladinian (Middle Triassic)



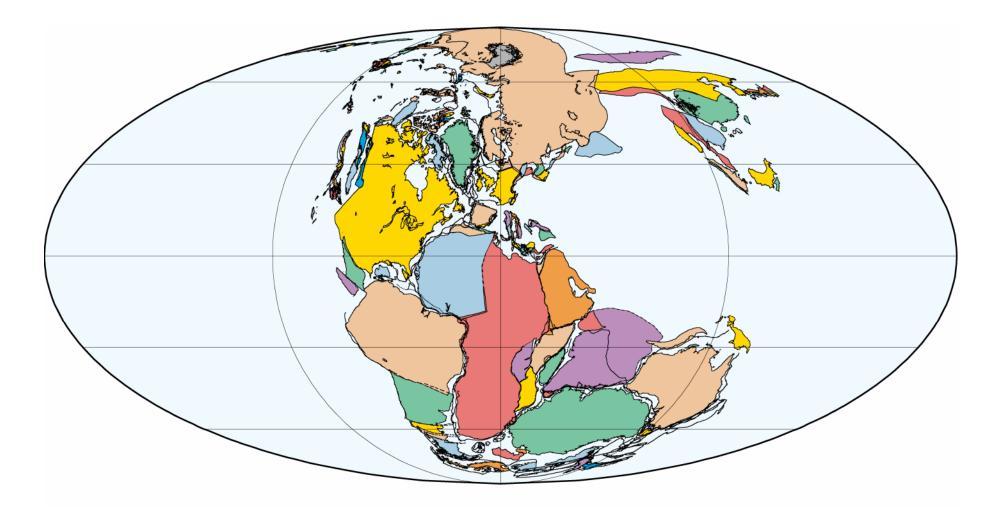
220 Ma Early Norian (Late Triassic)



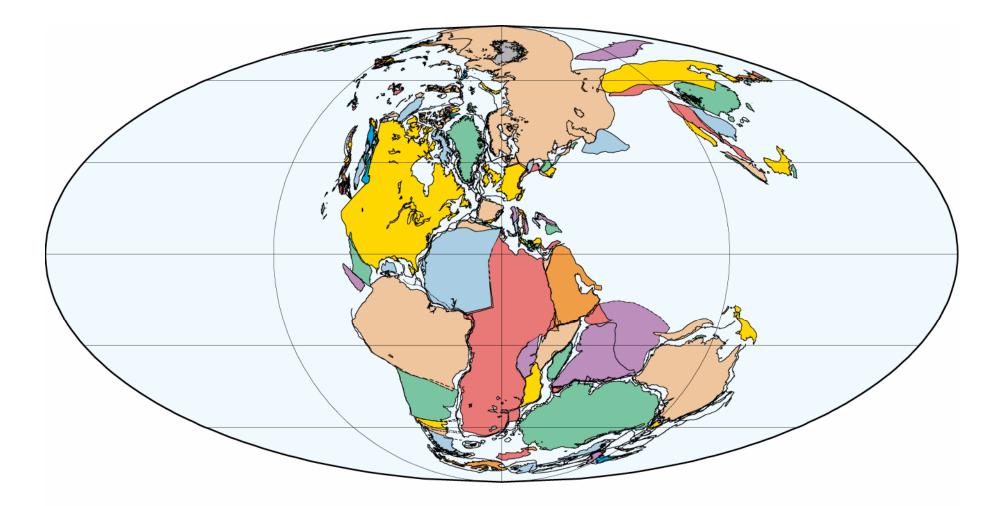
210 Ma Late Norian (Late Triassic)



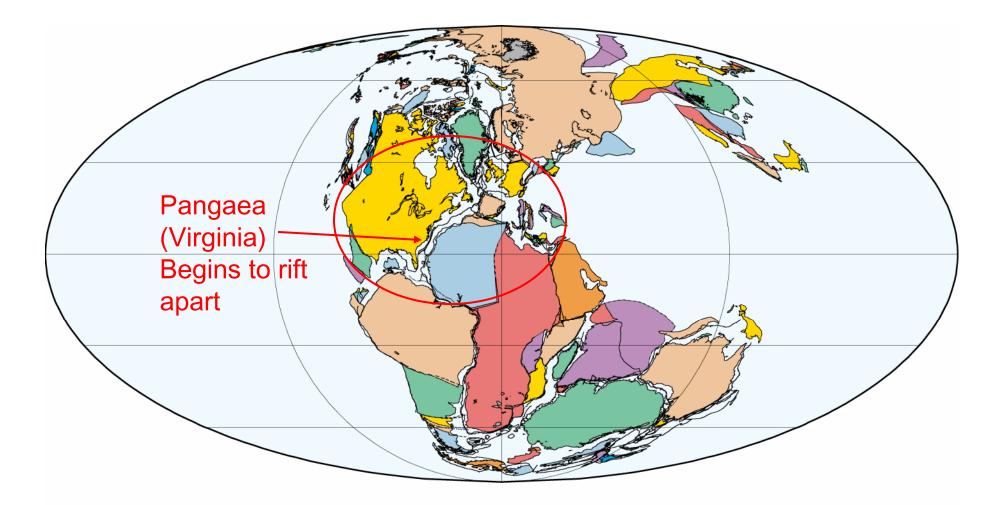
200 Ma Sinemurian (Early Jurassic)



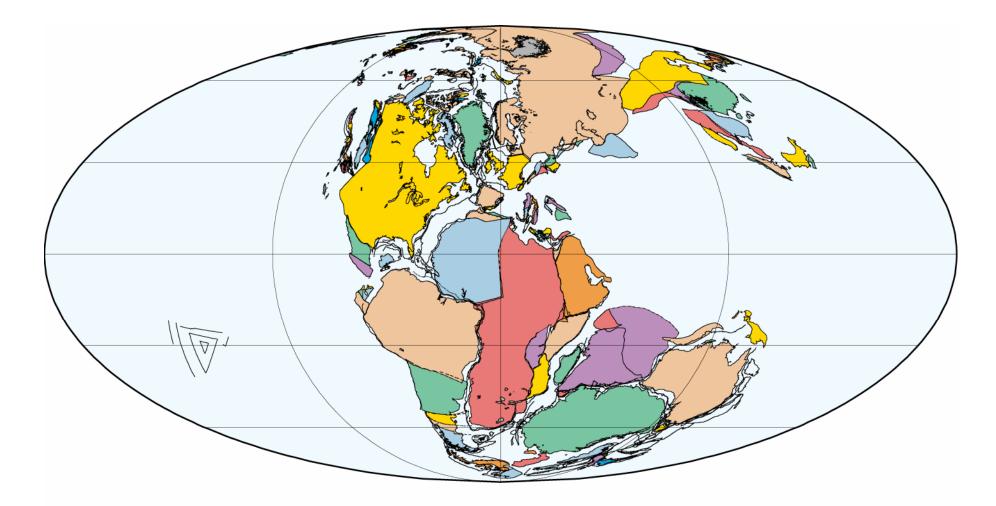
190 Ma Pliensbachian (Early Jurassic)



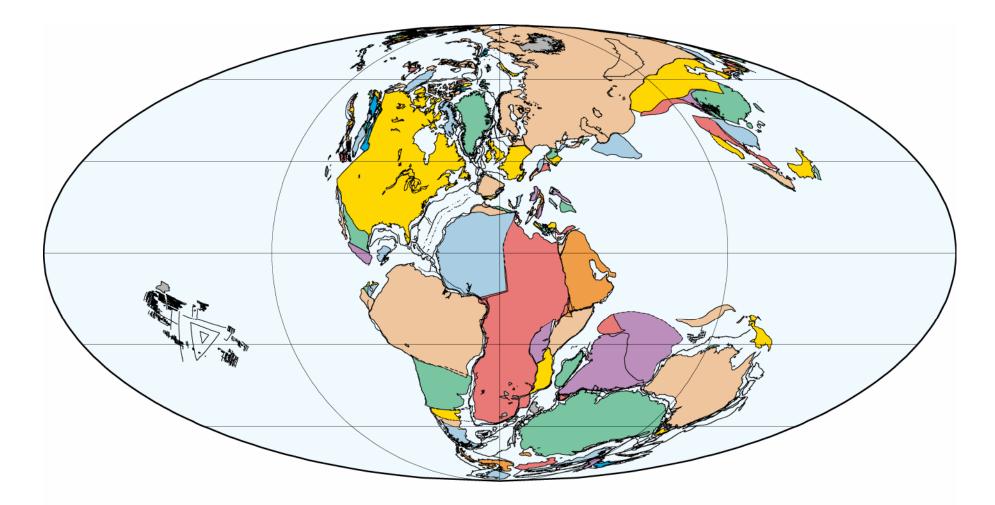
180 Ma Aalenian (Middle Jurassic)



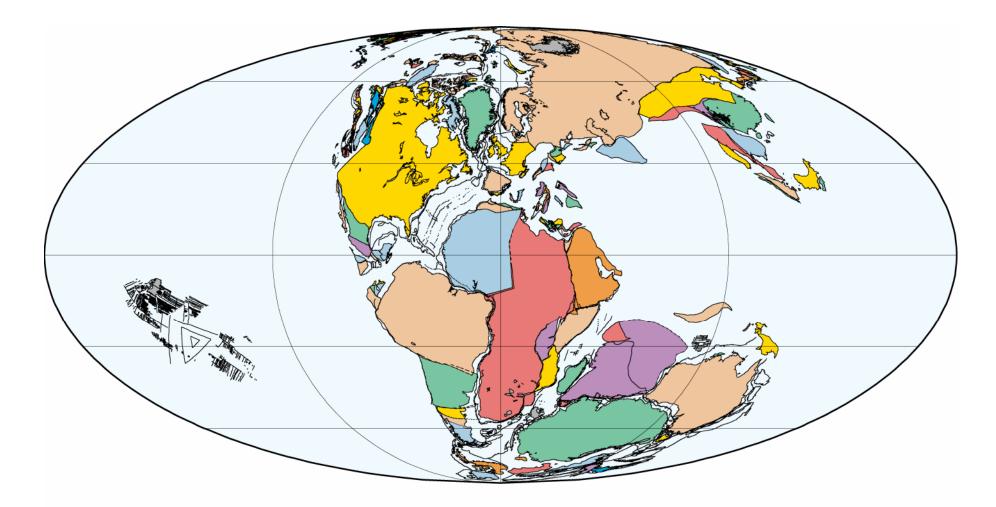
170 Ma Bajocian (Middle Jurassic)



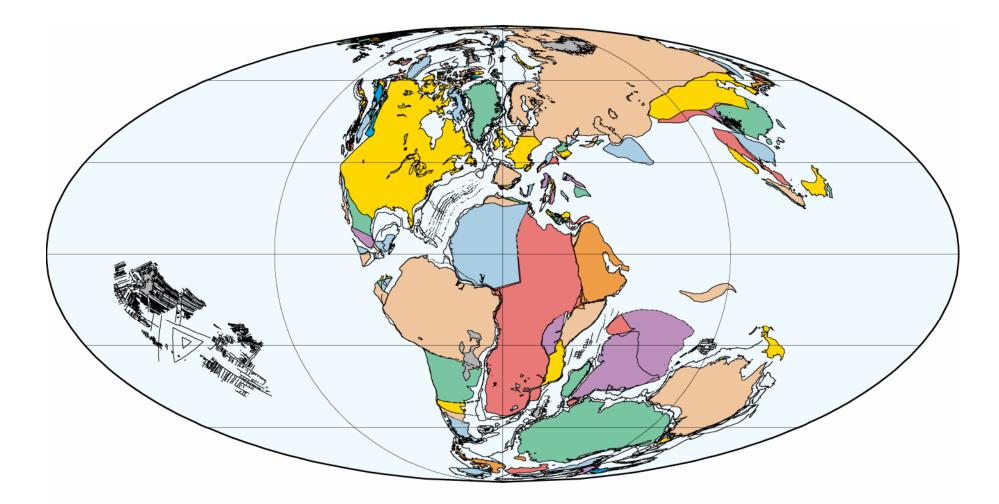
160 Ma Callovian (Middle Jurassic)



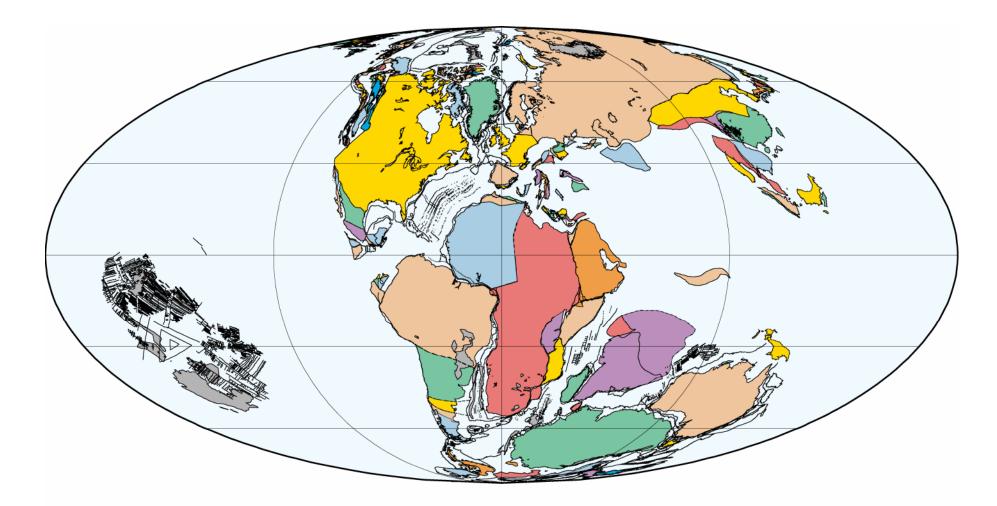
150 Ma Volgian (Late Jurassic)



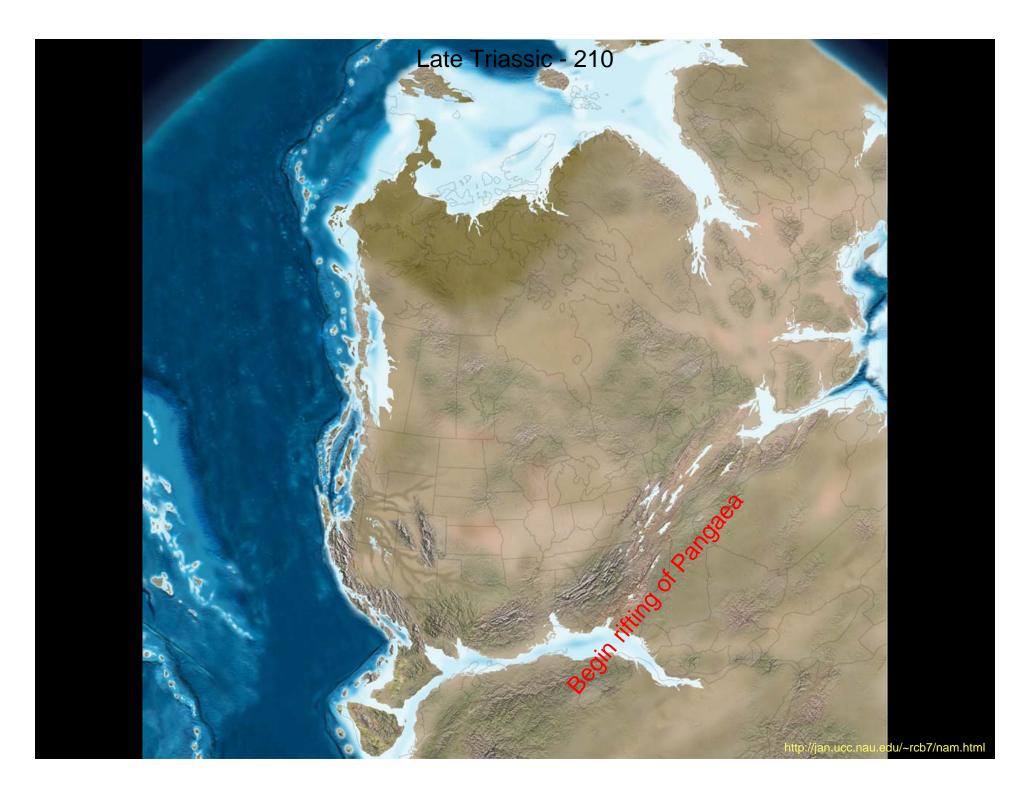
140 Ma Ryazanian (Early Cretaceous)

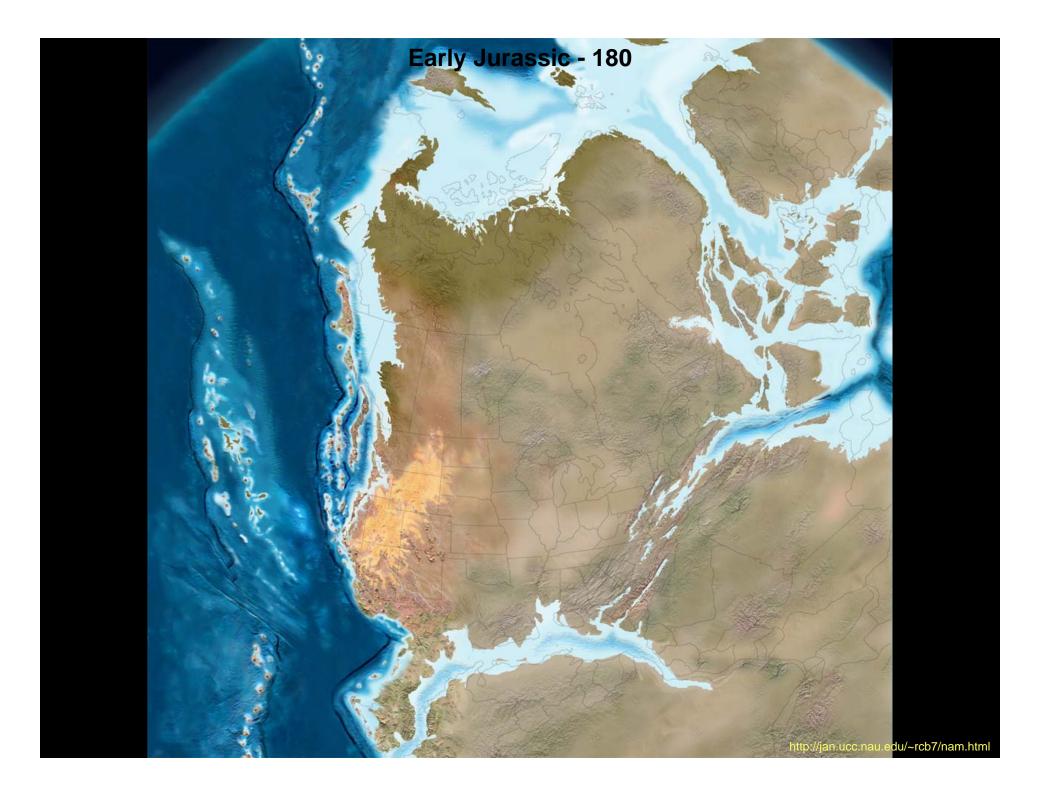


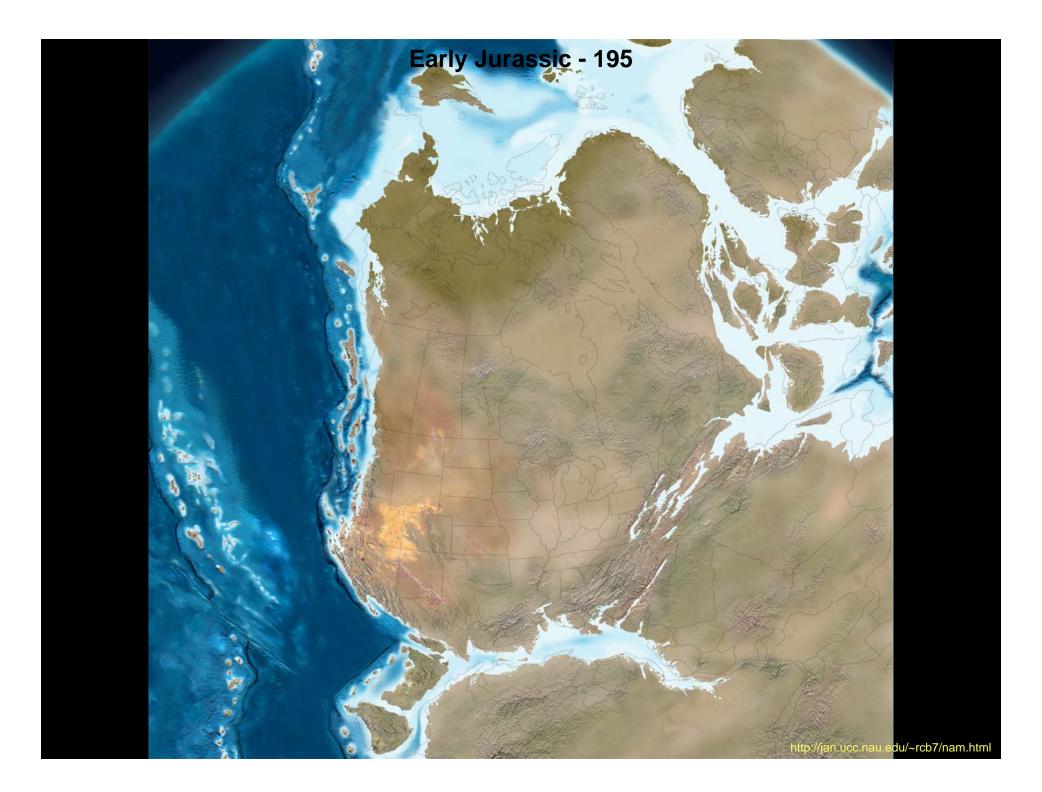
130 Ma Hauterivian (Early Cretaceous)

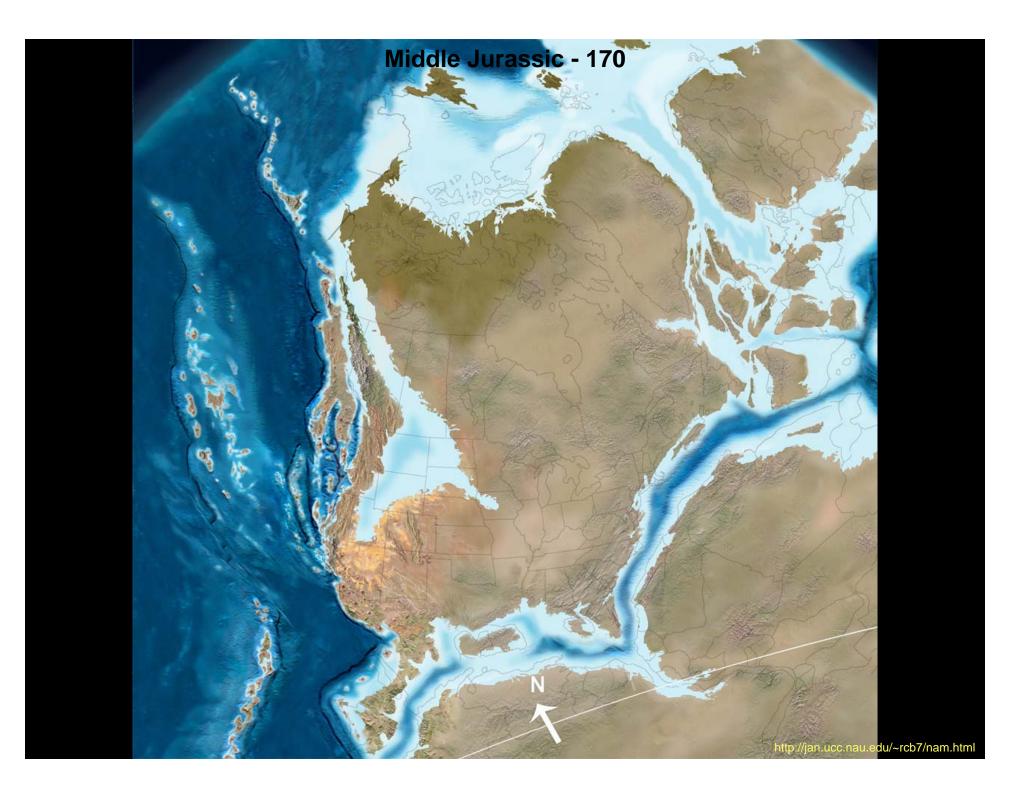


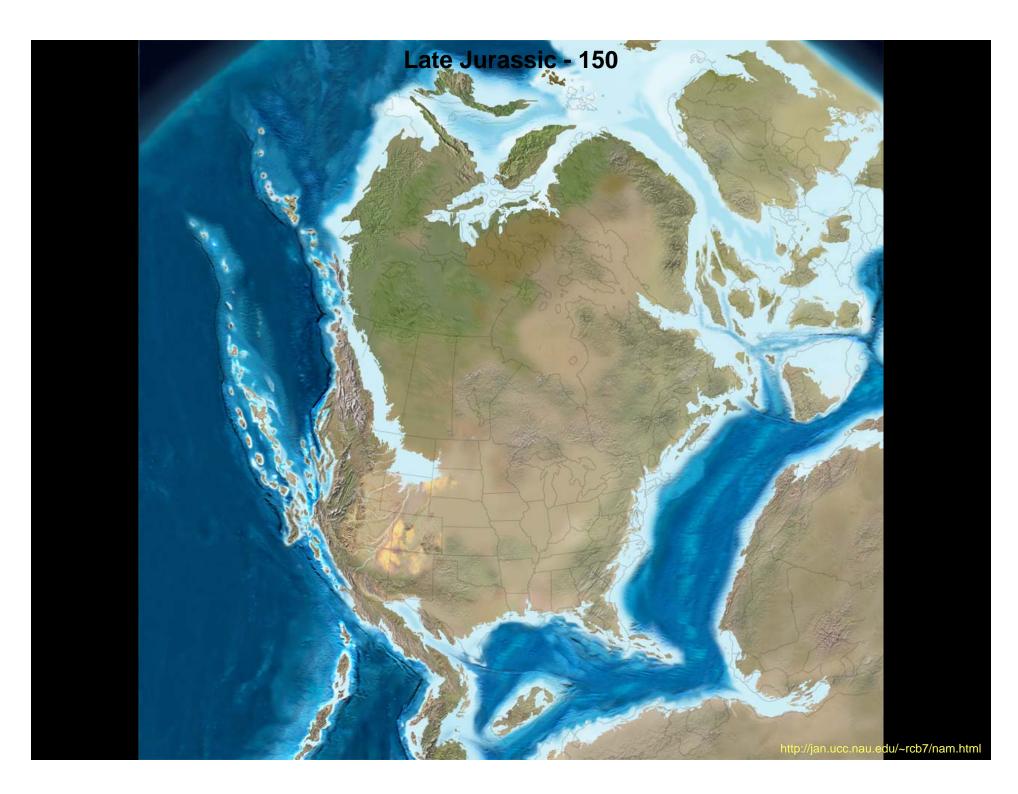
120 Ma Aptian (Early Cretaceous)

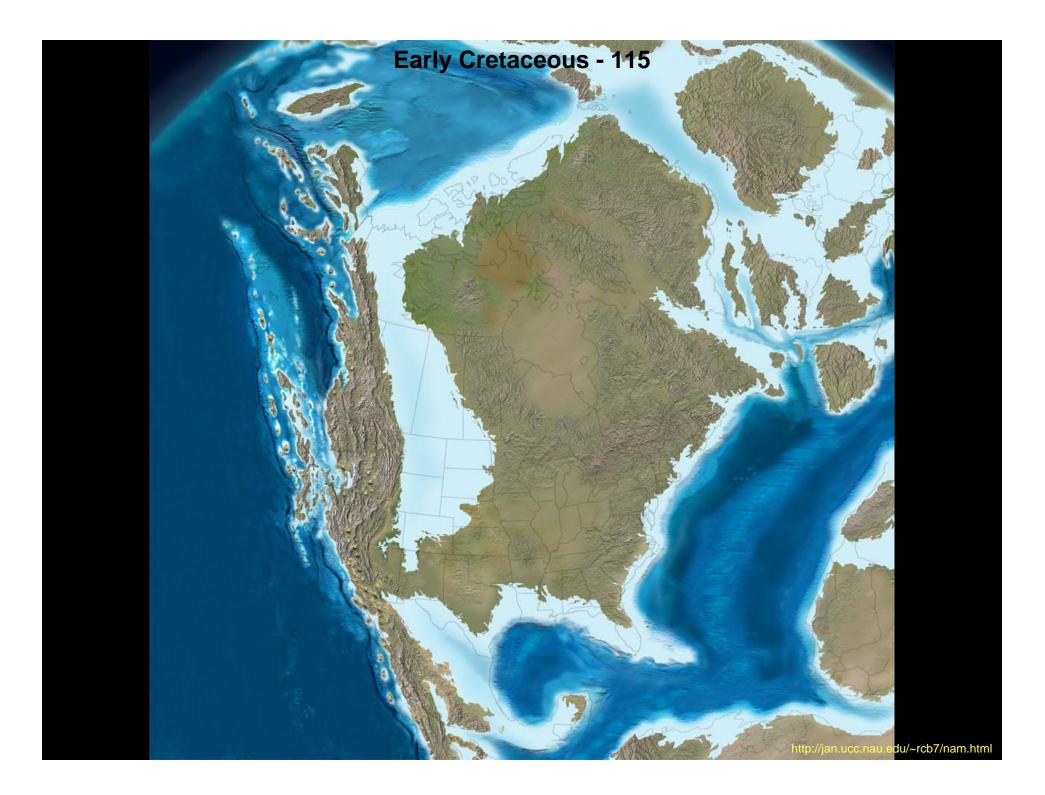


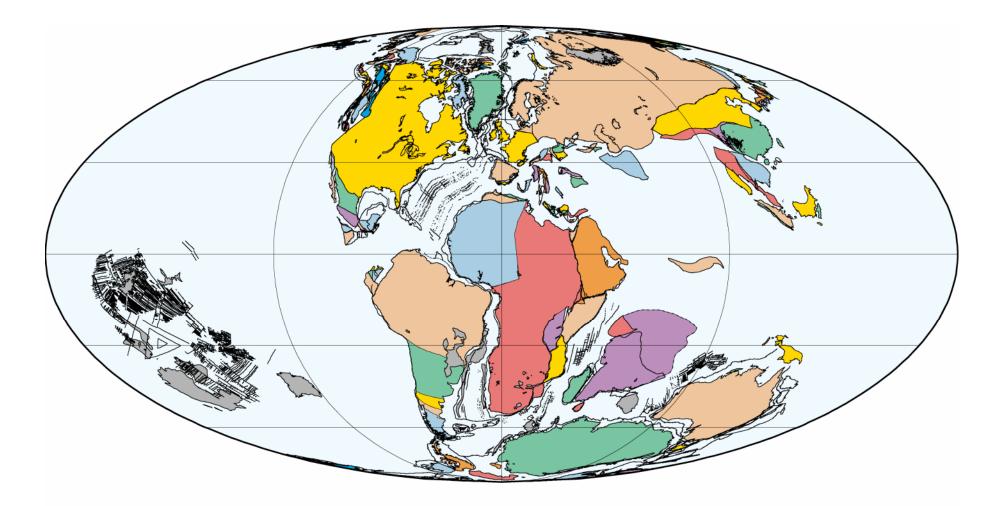




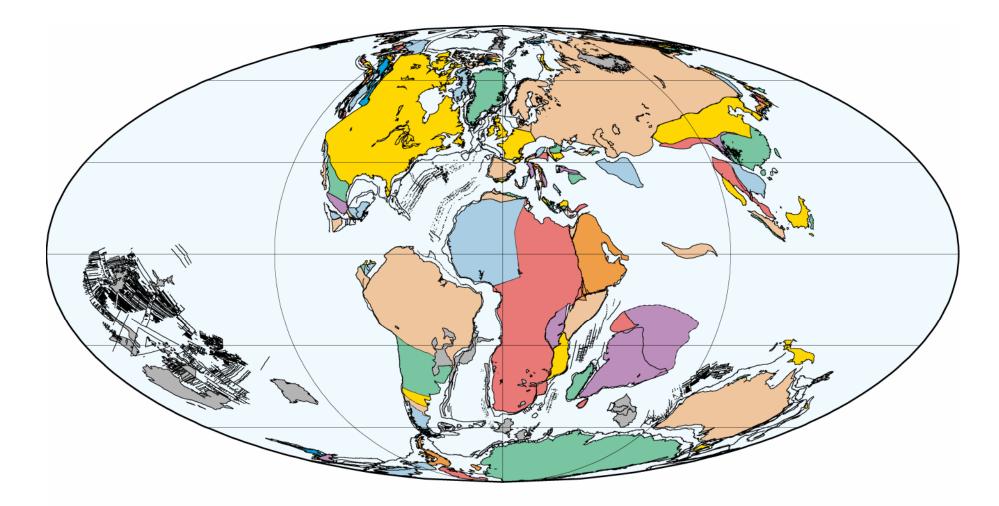




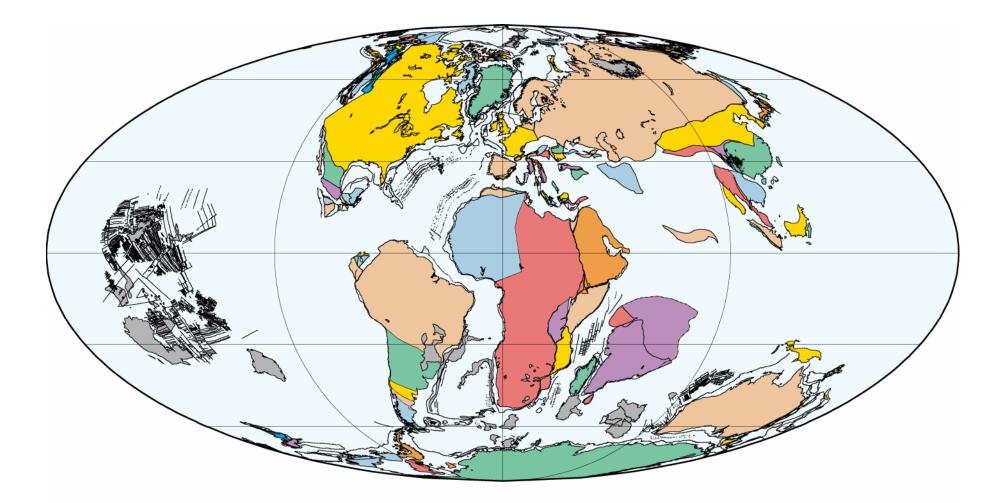




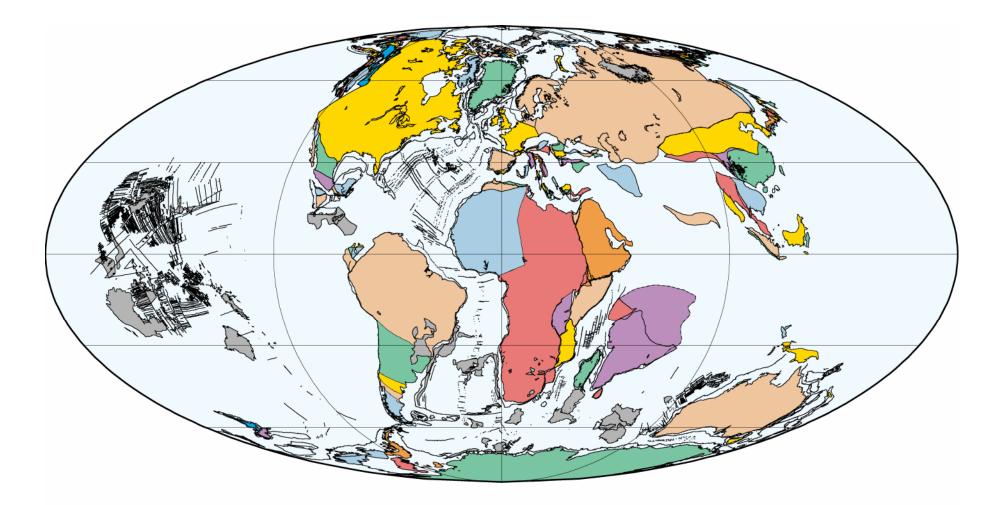
110 Ma Early Albian (Early Cretaceous)



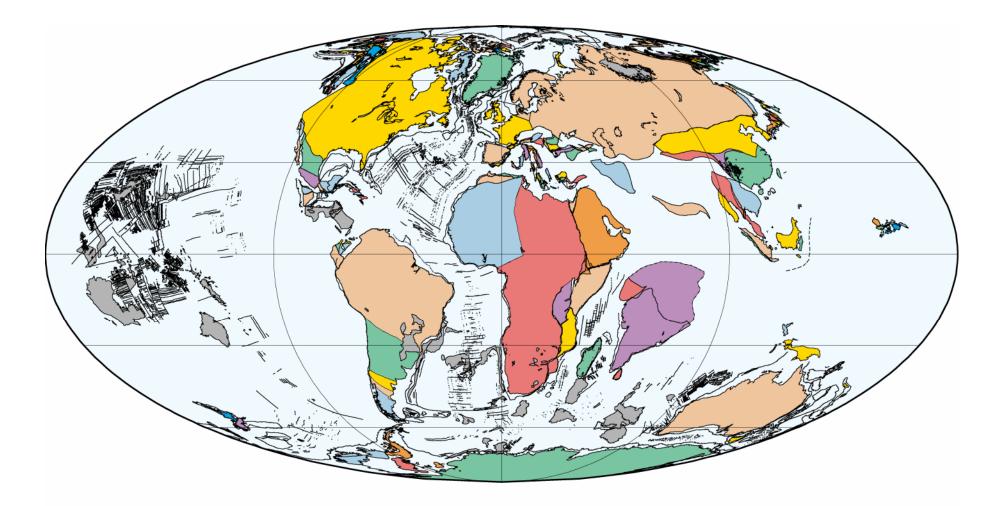
100 Ma Late Albian (Early Cretaceous)



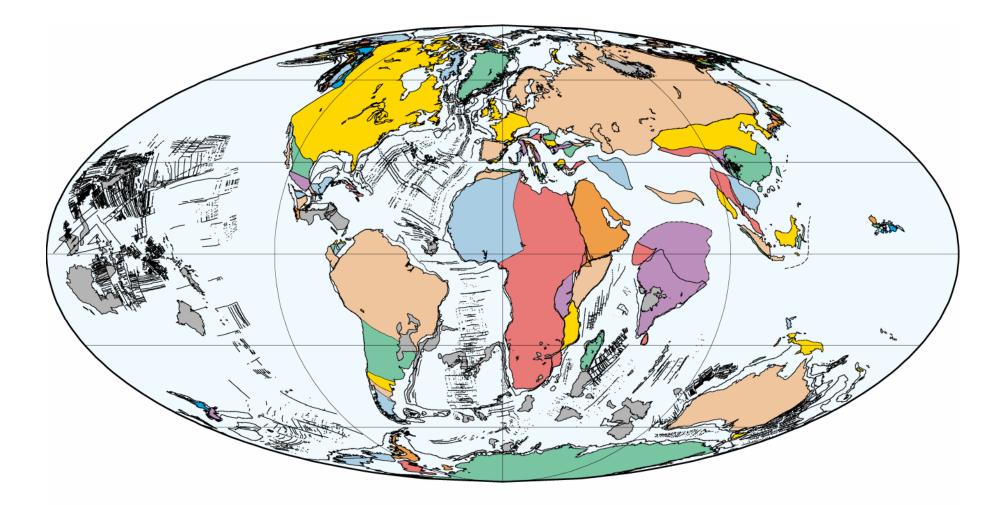
90 Ma Turonian (Late Cretaceous)



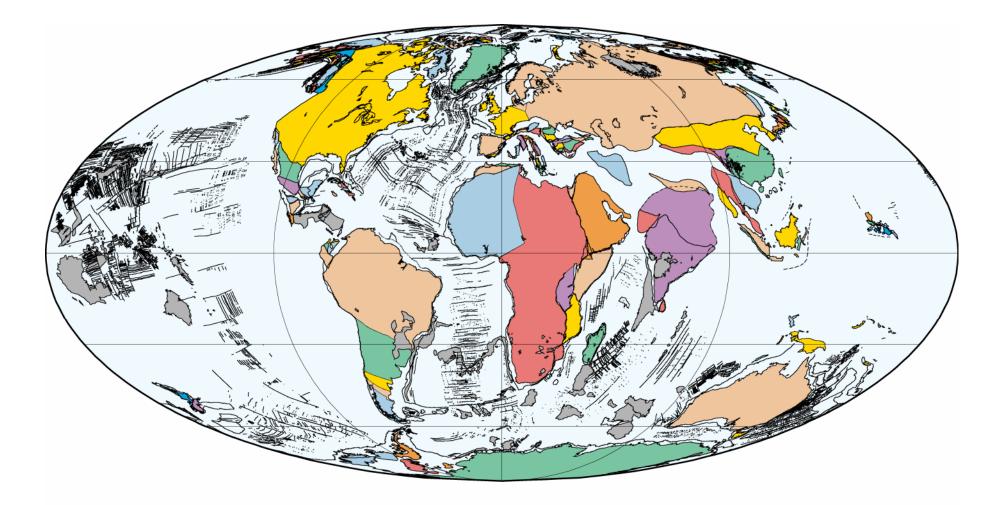
80 Ma Campanian (Late Cretaceous)



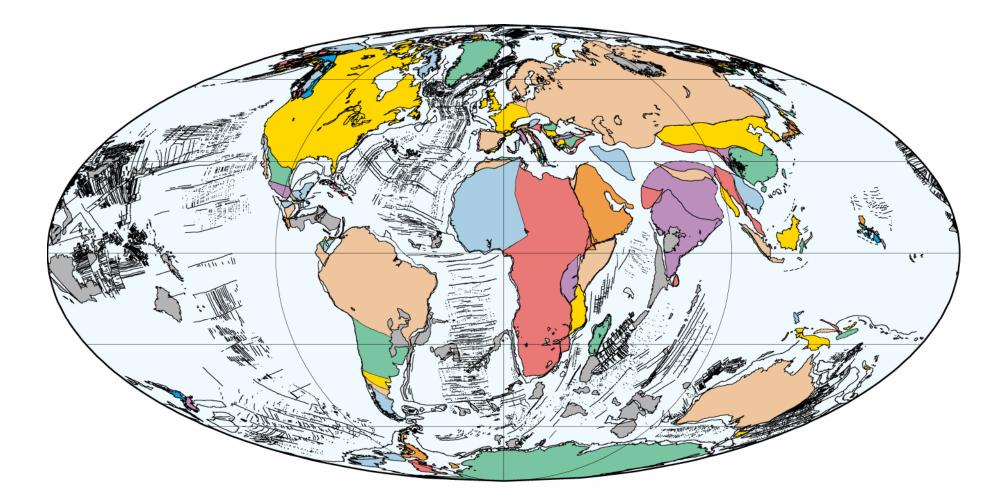
70 Ma Maastrichtian (Late Cretaceous)



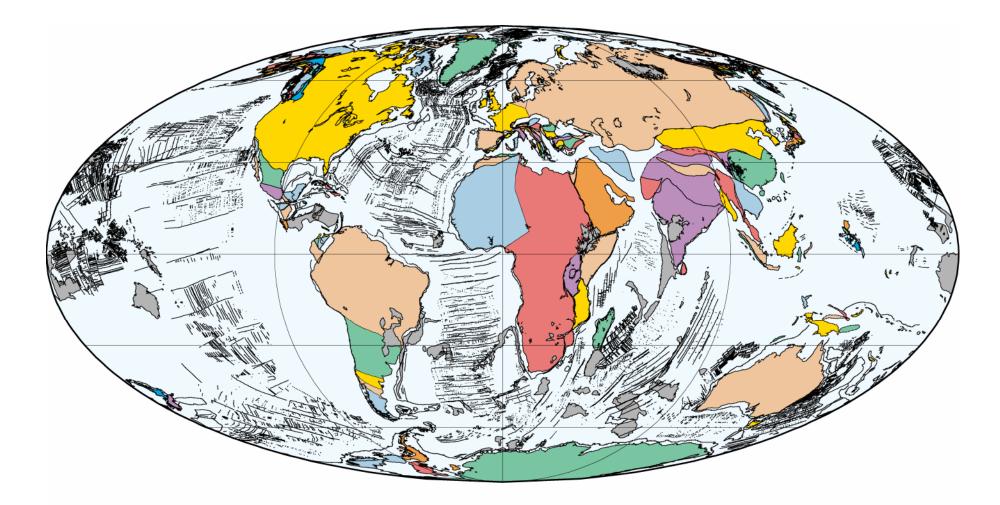
60 Ma Late Paleocene



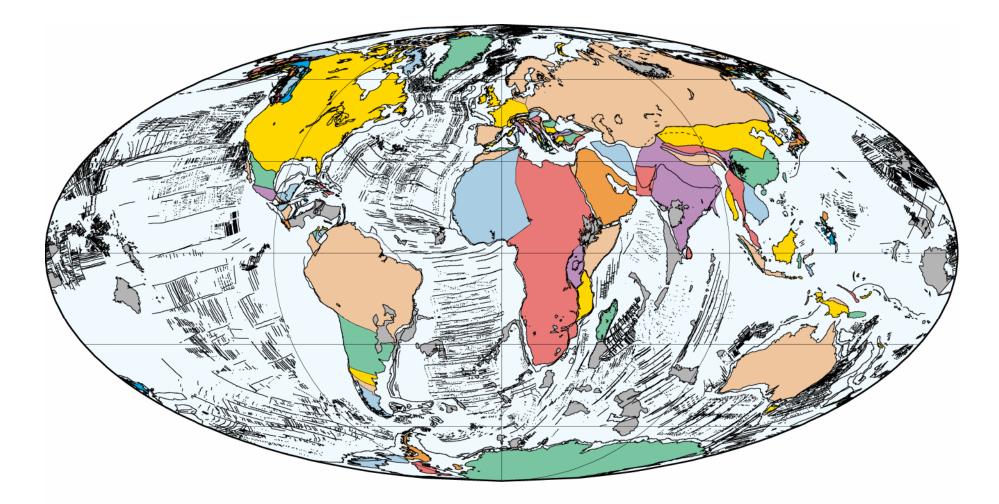
50 Ma Early Eocene



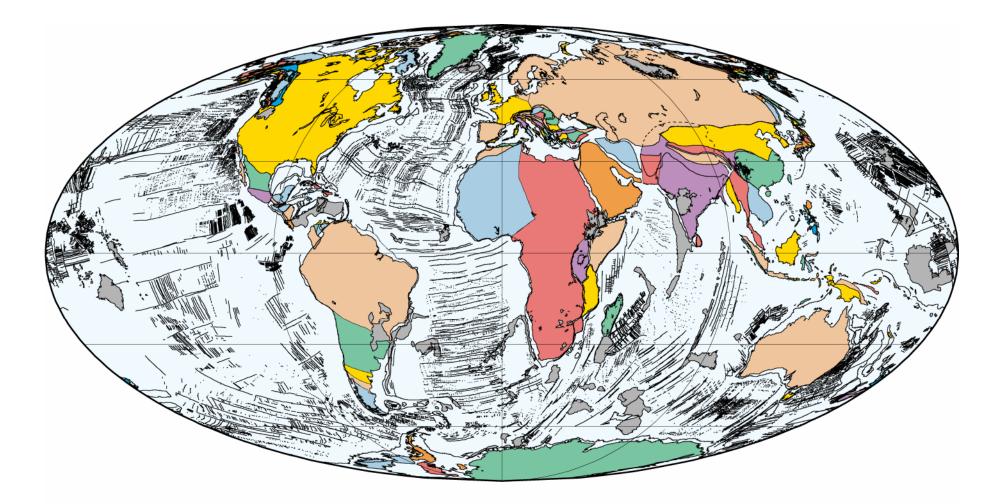
40 Ma Middle Eocene



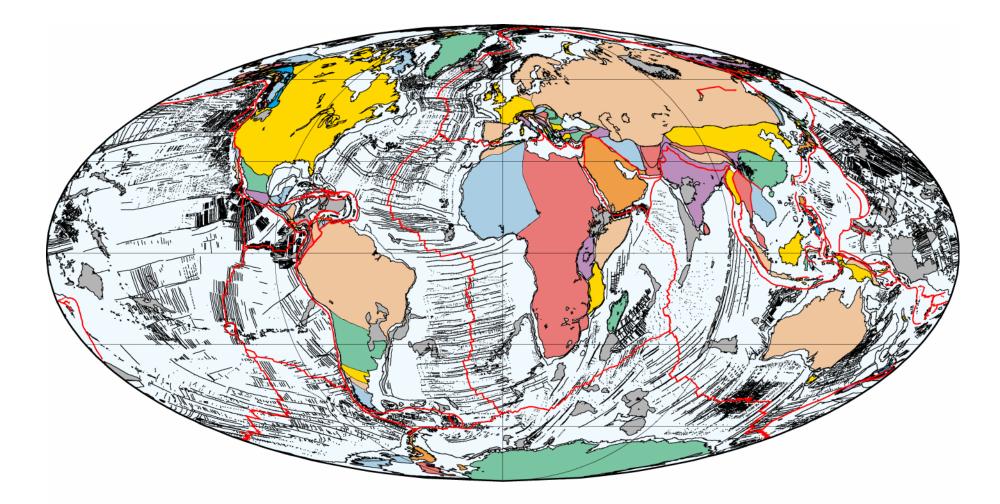
30 Ma Early Oligocene



20 Ma Early Miocene



10 Ma Late Miocene



0 Ma Present Day

